Midair Collisions Enhance Saltation

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Here we address the old question in aeolian particle transport about the role of midair collisions. We find that, surprisingly, these collisions do enhance the overall flux substantially. The effect depends strongly on restitution coefficient and wind speed. We can explain this observation as a consequence of a soft bed of grains which floats above the ground and reflects the highest flying particles. We make the unexpected observation that the flux is maximized at an intermediate restitution coefficient of about 0.7, which is comparable to values experimentally measured for collisions between sand grains.

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Would a sandstorm be stronger if the sand grains in air did not collide against each other? This question has puzzled practitioners and theoreticians alike in the past. Models for aeolian sand flux [1–6] become certainly much simpler if such midair collisions are neglected, but does this approximation underestimate or overestimate the value of the saturated flux? As opposed to experiments, the direct computer simulation of saltation, the main aeolian transport process, offers the possibility of switching on or off the collisions between particles or of modifying the collision parameters, such as the coefficient of restitution. This allows us, for the first time, to precisely determine the role of midair collisions during saltation.

We discover that midair collisions are the key ingredient for understanding the relation between different concepts such as the splash [7,8], the soft bed [9,10], and the distinction between saltons and reptons [11,12]. During saltation, particles are ejected from the granular bed in a splash, produced by the impact of fast particles, so-called saltons (yellow trajectory in Fig. 1). These saltons must have the necessary kinetic energy to assure that, despite the substantial dissipation in the ground, some ejected particles can fly sufficiently high. After all, only high fliers can acquire sufficient acceleration to again become saltons because the wind velocity at the ground is zero, increasing logarithmically with height. Our detailed study reveals the following picture: Because of surface irregularities, a splash produces three types of moving particles (see Fig. 1): Many (green) creepers that do not leave the bed, many (red) leapers making small jumps, remaining in regions of small wind velocities, thereby not being able to produce a new splash, and very few saltons (yellow) which fly higher up. Only saltons sustain saltation. Both creepers and leapers (reptons) contribute considerably to the sand flux, but play a very different role in what follows.

Can one really make such a sharp distinction between leapers and saltons? After all, the collision process is

random and should yield a continuous distribution of ejection velocities. The experimental observation of single impacts on granular packings shows a bimodal splash distribution [13], reproduced numerically by Anderson and Haff [14], exhibiting a broad spectrum of slower particles and a small peak of faster ones. The analysis of these observations led Andreotti [11] to coin the terms saltons and reptons. However, so far there is no experimental evidence that the same velocity distribution in steady state saltation is also bimodal. What makes these two types of particles different are the midair collisions. Figure 2 shows the trajectory of a typical salton simulated with the method described below in 3D: It makes several jumps without touching the ground, rebouncing upwards each time due to a collision with a leaper. The probability for such a collision is reasonably high because the leapers form a rather dense layer, which is precisely the soft bed described earlier [10]. Consequently, the salton stays longer in areas of strong wind and less time close to the ground, where the wind drag acceleration is weaker. This explains why the saltons acquire so much energy and can



FIG. 1 (color online). Typical splash obtained numerically with our model in 3D. The impinging particle (yellow) bounces after ejecting other particles from the bed (red and green). While the green particles essentially move only on the ground, the red ones are lifted and dragged by the wind.



FIG. 2 (color online). Simulated trajectory of a salton in 3D. The (yellow, upper trajectory) salton is kept in the air by colliding against (red, lower trajectories) particles from the soft bed. For clarity, only the relevant particle trajectories are shown and the ground is schematically represented by a flat plane.

sustain the saltation process. Summarizing, midair collisions in the soft bed is the crucial mechanism that differentiates the saltons and doubles the saturated flux as we will show here.

The discrete elements method, which is based on particle-particle and particle-wind interactions, allows us to quantitatively study the aeolian transport [15]. It has been used to calculate the onset of saltation and to confirm the existence of a jump in the total flux in 2D [16]. Here we apply a 3D scheme to investigate the role of midair collisions.

Every sand grain is represented by a hard sphere of average diameter D_{mean} . Gravity acts in vertical direction (y direction) and a logarithmic wind velocity profile u(y) is imposed in the horizontal direction (x direction),

$$u(y) = \frac{u_*}{\kappa} \ln \frac{y - h_0}{y_0},$$
 (1)

where $y_0 = D_{\text{mean}}/30$ is the roughness of the bed, h_0 the bed height, $\kappa = 0.4$ the von Kármán constant, and u_* the wind shear velocity. The Shields number, defined as

$$\theta = \frac{u_*^2}{(s-1)gD_{\text{mean}}},\tag{2}$$

is the pertinent dimensionless parameter that controls the wind velocity, where $s = \rho_s / \rho_w$ is the ratio between the grain and fluid density, and g the norm of the gravitational acceleration. The feedback procedure that extracts the momentum from the wind due to the acceleration of the grains is explained in the Supplemental Material [17].

We define midair collisions as those for which both particles have their center of mass above h_0 . Collisions with bed particles occur when the center of mass of at least one particle is below h_0 . We study the effects of the restitution coefficient e on midair collisions for a fixed restitution coefficient $e_{bed} = 0.7$ for collisions with the particle bed. In the simulations without midair collisions, above h_0 the collisions are neglected; i.e., the particles are transparent to each other. Further information can be found in the Supplemental Material [17].

We consider particles of average diameter $D_{\text{mean}} = 2 \times 10^{-4}$ m, size dispersion $\sigma_D = 0.15D_{\text{mean}}$, and density $\rho_s = 2650 \text{ kg/m}^3$, in a three-dimensional wind channel of dimension $(700 \times 50 \times 7.5)D_{\text{mean}}^3$ with a reflective upper boundary, placed sufficiently high to avoid any collision against it, and periodic boundaries in the other directions. For the fluid density, we chose $\rho_w = 1.174 \text{ kg/m}^3$. In fact, no collision with the upper boundary thas occurred in our simulations. The lower boundary at y = 0, representing the deep ground, is strongly dissipative with a fixed restitution coefficient of $e_w = 0.5$. We consider a bed of 12 particle layers to suppress the reflection of shock waves on the lower boundary due to finite depth [18–20]. In the beginning of the simulation, a few particles are dropped at random positions to trigger saltation.

The dimensionless flux in the direction of wind is defined as

$$q = \frac{1}{DA} \sum_{i}^{N} m_i v_i^x, \qquad (3)$$

where $A = (50 \times 7.5)D_{\text{mean}}^2$ is the area of the bottom of the channel, v_i^x and m_i are, respectively, the velocity along the horizontal direction and the mass of the particle *i*, and $D = \rho_s \sqrt{(s-1)gD_{\text{mean}}^3}$ is a normalization constant. The saturated flux, which is the average flux in the stationary state, does not change with the number of particles. The granular temperature, defined as

$$T(y) = \frac{1}{3N(y)} \sum_{i}^{N(y)} m_i (\mathbf{v}_i - \mathbf{v}_m(y))^2,$$
(4)

quantifies the fluctuations around the mean velocity $\mathbf{v}_m(y) = 1/N(y) \sum_{i}^{N(y)} \mathbf{v}_i$.

The wind channel is divided along the y direction in horizontal slices $(2.5 \times 50 \times 7.5)D_{\text{mean}}^3$ to calculate the profiles of particle concentration, average particle velocity, alignment, and granular temperature. The particle concentration is the ratio between the volume of particles and the total volume. The flux profile is obtained for each slice from the product of the concentration and the average particle velocity.

Figure 3(b) shows the dependence of the saturated flux on *e* for different θ . For $\theta > 0.44$, the saturated flux is higher for e = 0.75 than for e = 1.0. The maximum at e = 0.75 increases substantially with θ ; i.e., the stronger the wind, the higher the peak. This is also confirmed by the flux profile in Fig. 4(b), where the area below the curve, which corresponds to the total flux, is larger in the presence of dissipation. Interestingly, this optimal *e* is comparable to the values experimentally measured for collisions between quartz grains [21]. Below we will explain this maximum as the result of competition between the loss of alignment between trajectories and mounting uplift of particles with increasing *e*.



FIG. 3 (color online). (a) Concentration profile of particles as a function of the height y for $\theta = 0.44$ in the absence (red, lower curve) and the presence (yellow, upper curve) of midair collisions (e = 0.7). (b) The relation between the saturated flux and restitution coefficient, for four different θ exhibits a peak around e = 0.75 for higher shear velocities. The horizontal dashed lines show the flux without midair collisions.

In the absence of midair collisions, all particles are reptons and follow stretched parabolic trajectories [1]. Saltons emerge with midair collisions. They are located at higher positions where the wind is stronger (Fig. 2) and spend less time close to the ground, where wind drag acceleration is weaker. Thus, they are much faster and contribute much more to the overall flux. The number of collisions and, consequently, the flying time of a salton strongly depend on the concentration of leapers. Huang et al. have also recognized that midair collisions might sustain grains above the ground reducing the frequency of collisions with the bed [22–24]. However, in contrast to our observation, they hypothesized that such a decrease would reduce the mass transport. Indeed, we confirmed that the mass flux of the no-collision case is below the one of the collisional case (e = 0.7) even for small θ close to the transport threshold. Figure 3(a) compares the concentration profiles for $\theta = 0.44$ with and without midair collisions. The yellow and red (dark) curves are in good agreement with the ones obtained by Jenkins and co-workers with kinetic theory with [25] and without [26] midair collisions, respectively.

Simultaneously, another important mechanism is activated. While, without dissipation, particles rebound randomly in air, dissipation tends to align the trajectory of colliding particles [27]. This is expected to occur in the direction of wind. We define particle alignment as the variance of the velocity angle with respect to wind direction, i.e., $\sigma_{\alpha}^2 = \langle \alpha^2 \rangle - \langle \alpha \rangle^2$ with $\alpha = \arctan(v_y/v_x)$. A larger variance corresponds to a weaker alignment. To investigate this, we measure, for the dissipative (e = 0.7) and the conservative (e = 1.0) cases, the granular temperature T(y) and the inverse of the temperature anisotropy $a(y) = T_y/T_x$ with $T_y = \langle v_y^2 \rangle$ and $T_x = \langle (v_x - \langle v_x \rangle)^2 \rangle$. The granular temperature in Fig. 4(a) displays at every



FIG. 4 (color online). Temperature (a) and flux (b) profiles for e = 0.7, 1.0, and without midair collisions, for $\theta = 0.44$. At every height in (a), the granular temperature for e = 1.0 is larger than e = 0.7. The flux profiles in (b) confirm the higher flux for e = 0.7. The flux profile is defined at each height as the product of the concentration and the average velocity. (c) Dependence of the maximum temperature on θ .

height a larger temperature for e = 1.0 than for e = 0.7with a peak around $y = 50D_{\text{mean}}$. In the region of the soft bed, the granular temperature profile is an increasing function of height. This result is in line with what was observed using kinetic theory [25]. Figure 4(c) shows the dependence of the maximum temperature on θ for both cases. The maximum temperature grows with $\theta^{1/2}$ (linear in u_*) for e = 0.7 and increases linearly with θ (quadratically with u_*) for e = 1.0, being always larger in the conservative case. This higher temperature could be due to a larger dispersion either in magnitude or direction of particle velocities. To distinguish these two possibilities we plot, in Fig. 5(a), the variance of the velocity angle σ_{α}^2 and the inverse of temperature anisotropy a(y). The plots confirm that a lower temperature in the inelastic case (when compared with the elastic one) corresponds to a lower dispersion in particle velocities, i.e., higher alignment. Clearly, the alignment of trajectories, which tends to enhance the overall flux, decreases with e.

In the absence of midair collisions, the distribution of maximum heights can be approximated by a Poisson process and, therefore, is well described by an exponential decay, as shown in Fig. 5(b). With midair collisions, however, this distribution can only be described by the superposition of two exponentials: the first one corresponds to leapers in the simulation without midair collisions, while the second one corresponds to saltons, which typically fly above the leapers. We can now quantitatively define the upper bound of soft bed as the height above which saltons are in a majority. From the trajectories of individual particles, we can also relate the flying time to the maximum height [inset, Fig. 5(b)]; we observe that saltons stay much longer in the air than leapers.

Three major types of transport contribute to the total flux in saltation: sheet flow, transport of leapers (in the soft



FIG. 5 (color online). (a) The angle variance between particle trajectories for e = 0.7, e = 1.0, and without midair collisions, for $\theta = 0.44$. In the inset, the profile of the inverse of the temperature anisotropy. Also for $\theta = 0.44$, (b) shows the probability distribution for the maximum height, with (black circles) and without (blue triangles) midair collisions. The results in the absence of midair collisions can be fitted by the dashed (yellow) curve, given by $[0.041 \exp(-0.052x)]$, and include only leapers. The case with midair collisions is fitted by the solid (red) curve which is a superposition of the contribution of leapers and the one of the saltons [dash-double-dotted (purple) curve], given by $[0.6365 \exp(-0.15x) + 0.02 \exp(-0.026x)]$. In the inset, we show the relation between the flying time and the maximum height, where the horizontal dashed line delimits the soft bed.

bed), and of saltons (above the soft bed). Once the limits of the soft bed are identified, we can compute for each θ the contribution of each mechanism to the total flux, as shown in Fig. 6. The sheet flow, computed from the mass transport in the particle bed ($y < h_0$), results from the wind shear stress and the creep of particles on the surface. The contributions of leapers and saltons are obtained from the fluxes in and above the soft bed, respectively. As observed in the figure, the relative contribution of saltons and leapers significantly increases with θ in comparison to the sheet



FIG. 6 (color online). Contribution to the total saturated flux by sheet flow (lower curve), leapers in the soft bed (middle curve), and saltons above the soft bed (upper curve) for e = 0.7 as function of θ . The green area (lower region) shows the sheet flow in the particle bed $y < h_0$.

flow. Figure 6 also shows that the saltons contribute the most to the total flux.

For the saturated flux, we also reproduced the discontinuity at the onset of saltation reported for 2D [16]. We also studied the impact of midair collisions in 2D. The decrease in the spatial dimension enhances the relevance of midair collisions, and the consequences discussed here are even more pronounced. However, the curves for the saturated flux in 2D and 3D overlap if θ in 2D is rescaled by an appropriate factor that takes into account the system width.

Summarizing, the contribution of midair collisions cannot be neglected in saltation, as it would underestimate the mass flux. The saltons contribute significantly to the flux enhancement by acquiring large momentum from the wind and using it to eject more particles through splashes. We provide a new picture of aeolian saltation, in which the competition between uplift due to midair collisions and alignment due to inelasticity, optimizes the mass transport in the presence of dissipative collisions. Interestingly, the position of the maximum corresponds to the restitution coefficient typically observed in granular collisions. Berger et al. have shown a nontrivial dependence of the erosion rate on the restitution coefficient of midair collisions during a lunar landing in the absence of saltation [28]. As a follow-up, continuum models of aeolian transport would profit from a systematic analysis of the dynamics of particle ejection from the surface in the presence of a soft bed.

Our results are crucial for future studies since they provide qualitative and quantitative information about the influence of midair collisions, which should be considered in modeling. Additional work might include the aerodynamic lift of the particles, turbulent wind speed fluctuations, and the electrostatic interaction between particles. Further possibilities are the study of the role of collisions in saltation under terrestrial desert wind, subaqueous, or Martian conditions. It is also still an open question why the maximum granular temperature scales linearly with the shear velocity in the inelastic and quadratically in the elastic case.

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- M. P. Almeida, J. S. Andrade, and H. J. Herrmann, Phys. Rev. Lett. 96, 018001 (2006).
- [2] M. P. Almeida, E. J. R. Parteli, J. S. Andrade, and H. J. Herrmann, Proc. Natl. Acad. Sci. U.S.A. 105, 6222 (2008).
- [3] B.T. Werner, Ph.D. thesis, Caltec, Pasadena, 1987.

- [4] I. K. McEwan and B. B. Willetts, J. Fluid Mech. 252, 99 (1993).
- [5] J.F. Kok and N.O. Renno, J. Geophys. Res. 114, 149 (2009).
- [6] T. Pähtz, J. Kok, and H. J. Herrmann, New J. Phys. 14, 043035 (2012).
- [7] R. A. Bagnold, *The Physics of Blown Sand and Desert Dunes* (Methuen, New York, 1941), p. 265.
- [8] J. F. Kok, E. R. Parteli, T. Michaels, and D. B. Karam, Rep. Prog. Phys. 75, 106901 (2012).
- [9] S. H. Williams, Ph.D. thesis, Arizona State University, 1987.
- [10] M. Soerensen and I. McEwan, Sedimentology 43, 65 (1996).
- [11] B. Andreotti, J. Fluid Mech. 510, 47 (2004).
- [12] M. Lämmel, D. Rings, and K. Kroy, New J. Phys. 14, 093037 (2012).
- [13] S. Mitha, M.Q. Tran, B.T. Werner, and P.K. Haff, Acta Mech. 63, 267 (1986).
- [14] R.S. Anderson and P.K. Haff, Science 241, 820 (1988).
- [15] M. Griebel, S. Knapek, and G. Zumbusch, Numerical Simulation in Molecular Dynamics: Numerics, Algorithms, Parallelization, Applications (Springer, New York, 2007).
- [16] M. V. Carneiro, T. Pähtz, and H. J. Herrmann, Phys. Rev. Lett. 107, 098001 (2011).

- [17] See Supplemental Material at http://link.aps.org/ supplemental/10.1103/PhysRevLett.111.058001 for more details about the model.
- [18] F. Rioual, A. Valance, and D. Bideau, Phys. Rev. E 62, 2450 (2000).
- [19] F. Rioual, Ph.D. thesis, University of Rennes 1, 2002.
- [20] F. Rioual, A. Valance, and D. Bideau, Europhys. Lett. 61, 194 (2003).
- [21] M. Banks, N. Bridges, and M. Benzit, in *Proceedings of the 36th Annual Lunar and Planetary Science Conference*, 2005, p. 2116.
- [22] Z. Dong, N. Huang, and X. Liu, J. Geophys. Res. 110, D24113 (2005).
- [23] S. Ren and N. Huang, Eur. Phys. J. E 33, 351 (2010).
- [24] N. Huang, Y. Zhang, and R. D'Adamo, J. Geophys. Res. 112, D08206 (2007).
- [25] J. M. Pasini and J. T. Jenkins, Phil. Trans. R. Soc. A 363, 1625 (2005).
- [26] J. T. Jenkins, I. Cantat, and A. Valance, Phys. Rev. E 82, 020301 (2010).
- [27] Y. L. Duparcmeur, H. Herrmann, and J. Troadec, J. Phys. I (France) 5, 1119 (1995).
- [28] K. J. Berger, A. Anand, P. T. Metzger, and C. M. Hrenya, Phys. Rev. E 87, 022205 (2013).