Quasistatic deformation of yield stress materials: Homogeneous or localized?

E. A. Jagla

Centro Atómico Bariloche, Instituto Balseiro, Comisión Nacional de Energía Atómica, CNEA, CONICET, UNCUYO, Avenida E. Bustillo 9500, R8402AGP San Carlos de Bariloche, Río Negro, Argentina

(Received 16 May 2023; accepted 5 September 2023; published 28 September 2023)

We analyze a mesoscopic model of a shear stress material with a three-dimensional slab geometry, under an external quasistatic deformation of a simple shear type. Relaxation is introduced in the model as a mechanism by which an unperturbed system achieves progressively mechanically more stable configurations. Although in all cases deformation occurs via localized plastic events (avalanches), we find qualitatively different behavior depending on the degree of relaxation in the model. For no or low relaxation, yielding is homogeneous in the sample, and even the largest avalanches become negligible in size compared with the system size (measured as the thickness of the slab L_z) when this is increased. On the contrary, for high relaxation, the deformation localizes in an almost two-dimensional region where all avalanches occur. Scaling analysis of the numerical results indicates that in this case, the linear size of the largest avalanches is comparable with L_z , even when this becomes very large. We correlate the two scenarios with a qualitative difference in the flow curve of the system in the two cases, which is monotonous in the first case and velocity weakening in the second case.

DOI: 10.1103/PhysRevE.108.034123

I. INTRODUCTION

Yield stress materials [1-3] are substances that share features of both solids and liquids. Below a well-defined applied shear stress, known as the critical stress σ_c , they deform elastically and reversibly. However, if this critical stress is exceeded, the material responds plastically, attaining a finite rate of deformation (or strain rate) noted as $\dot{\gamma}$. The dependence of $\dot{\gamma}$ on σ defines the flow curve of the material. In a large class of yield stress materials, the behavior of $\dot{\gamma}(\sigma)$ is continuous at σ_c : $\dot{\gamma}$ increases smoothly from zero when σ is increased beyond σ_c , typically as a power law of the form $\dot{\gamma} \sim (\sigma - \sigma_c)^{\beta}$ [4,5]. In other cases, there may be a discontinuous transition at σ_c , typically with some hysteresis [6–9]. The discontinuous case is also referred to as "reentrant," as it is typically obtained when the flow curve displays a negative slope region (Fig. 1) that is qualitatively similar to the one occurring in a liquid-gas transition in the van der Waals approximation. A discussion, in a unified framework, of the different phenomenology associated to these two possible scenarios is one of the aims of this work.

Deformation of yield stress materials proceeds via discrete plastic rearrangements generically called "avalanches" that reduce the shear stress locally, while increasing it in other parts of the sample via anisotropic elastic interactions. In a stress controlled experiment, if $\sigma < \sigma_c$, the material reaches—after a transient—a static equilibrium state in which the deformation rate is zero. On the other hand, if $\sigma > \sigma_c$, there are, at any moment, plastically active places in the sample that produce a finite average value of the deformation rate $\dot{\gamma}$.

Within the protocol of applying a fixed value of σ , the identification of isolated avalanches is not possible. However, individual avalanches can be identified and characterized by considering a strain driven experiment, in which the average strain in the sample γ is controlled. If the value of γ is

assumed to increase infinitely slowly with time (thus $\dot{\gamma} \rightarrow 0$), the stress in the sample increases continuously everywhere until there is a place where a mechanical instability occurs (typically a saddle-node bifurcation), and an avalanche is triggered. This avalanche produces a reduction of the stress in the system that can be taken as a measure of the size of the avalanche. The collection of points affected by the avalanche can also be determined, defining the avalanche spatial extent. Upon the further increase of γ , a sequence of avalanches will be observed and their statistical properties (such as size and duration distributions) can be studied.

The discrete nature of the deformation of yield stress materials has received a lot of attention in recent years, and in particular the properties of the avalanches responsible for the overall behavior have been studied in detail [10-13]. Yet, although the deformation of all yield stress materials can be described in terms of individual avalanches, it is crucial to understand the different effect that these avalanches may have in the behavior of macroscopic samples. There are cases in which the effect of individual avalanches becomes less and less detectable when system size increases. This corresponds to a case in which the behavior of a macroscopic sample can be characterized as "ductile." In other cases, the effect of individual avalanches continues to be detectable even if the sample size is increased more and more. We will refer to the behavior in such cases as "fragile"[14].

One of the main points we want to convey in this work is that the distinction between ductile and fragile behavior is intimately related to the underlying flow curve of the material being of the continuous or reentrant type. Another important difference between the two cases will emerge from the analysis: while ductile behavior is associated with avalanches that occur all across the sample (therefore producing, on average, a spatially uniform strain rate), the fragile case produces avalanches that tend to localize spatially in a quasi-two-



FIG. 1. Qualitative flow curve of a shear stress material with continuous (left) or discontinuous (right) yielding behavior. The dotted part of the flow curve in (b) is unstable and not experimentally observable, as it is typically replaced by some sort of generic Maxwell construction.

dimensional region where the strain rate remains finite, while it is essentially zero in the rest of the sample. This behavior can be described as the sample developing a "fault" in it, where all deformation occurs. In fact, a second idea that we want to convey is that materials with fragile behavior deform through the production of "earthquakes" that localize in a sort of "seismic fault." Beyond this localization effect, an important characteristic that will emerge is that individual earthquakes may have a macroscopically detectable effect, contrary to the case of avalanches in systems without relaxation. This will make a connection with the deeply established idea in the geophysical community that in order to sustain earthquakes, a material must have a friction law that is velocity weakening, i.e., a region in which the stress in the system decreases as a function of the strain rate [15]. This is precisely the situation with materials having a reentrant flow curve.

When materials with reentrant yielding are driven at a finite value of $\dot{\gamma}$, deformation localizes spatially in the form of a shear band [16-21]. In fact, a situation reminiscent of coexistence in a first-order phase transition occurs and there is a spatial separation between a nonflowing region and a flowing region in the form of a shear band in the system. The properties of shear bands have received much attention from the materials science community, as they are intimately related, for instance, to the failure mechanisms of materials such as metallic glasses [22-25]. The width of a shear band decreases as the global strain rate decreases, as a kind of Maxwell construction argument easily shows. In the limit of very low deformation rate $(\dot{\gamma} \rightarrow 0)$, the thinning of shear bands leads to the formation of a "fault" where earthquakelike avalanches occur. Yet, numerical models devised to study the statistical properties of earthquakes do not typically consider a bulk three-dimensional sample. Instead, it is usually assumed from the beginning that the deformation is localized in a two-dimensional fault, and therefore only this fault is modeled in detail, perhaps with the surrounding three-dimensional medium considered at a mean field level [26–28]. One standard approach is to model the two-dimensional fault as a collection of blocks that is joined by springs. The velocity weakening feature is incorporated either as a direct external ingredient (as in the Burridge-Knopoff model [26], in the form of the local sliding law of each block) or it emerges as a consequence of a more fundamental relaxation mechanism, such as the viscoelasticity model discussed in [29]. In either

case, the three dimensionality of the system is incorporated only through the stiffness of a driving spring acting on the effective two-dimensional fault. In addition, most simulations of two-dimensional faults consider the elastic interaction within the fault to occur among "nearest neighbors" only. Actually, in a real situation, this interaction is mediated by the bulk three-dimensional medium, and it is long range in an infinite sample.

In the present work, we intend to provide a comparative view of the cases of yield stress materials with continuous or reentrant behavior, and the relation to the earthquake phenomenology. We use a three-dimensional model, therefore not assuming any simplifying two-dimensional situation. We take a system of size $L \times L \times L_z$ (where L is supposed to be "very large"; see below), with periodic boundary conditions, and apply a simple shear deformation of the form $\mathbf{u} = \gamma z \hat{x}$, with γ defining the strain. The variation of L_{τ} allows us to perform a scaling analysis of the size of the observed avalanches. Note that L_z would eventually relate to the value of the driving spring that simplifying two-dimensional models use for earthquakes, and also controls the detailed form of the elastic interaction between different points of the fault. For each value of L_z , the value of L is chosen to be sufficiently large in such a way that the properties of the avalanches are not affected by a further increase of L. It is the value of L_7 that controls the "size effects" in the system.

The model (details in the next section) is driven through the quasistatic increase of the strain γ . This produces a sequence of avalanches that can be characterized in detail. In addition, relaxation is introduced through a parameter *R*, and the observed phenomenology crucially depends on the ratio $R/\dot{\gamma}$ [30]. In short, relaxation is a mechanism by which an unperturbed system achieves progressively mechanically more stable configurations.

Here is a brief summary of the main results that are obtained. When $R/\dot{\gamma} \rightarrow 0$, we reproduce well-known results corresponding to a system with a continuous yielding transition. In particular, we obtain a distribution of avalanche size of the form $P(S) \simeq S^{-\tau} \exp(-S/S_{\text{max}})$ with a value of $\tau \simeq 1.5$, and a cutoff avalanche size S_{max} that scales with L_z as $S_{\text{max}} \simeq L_z^{1.2}$. The actual observation of the avalanches reveals that although they display the spatial correlations expected from the symmetry of the elastic interaction, they appear all across the sample. Therefore, in the long run, the deformation of the system is uniform in this case.

For finite $R/\dot{\gamma}$, we observe the following trends:

(a) Avalanches become localized along the *z* axis, defining a quasi-two-dimensional region that can be called a "seismic fault."

(b) The distribution P(S) becomes broader, with S_{max} increasing with $R/\dot{\gamma}$ for a fixed L_z . The dependence of S_{max} with L_z displays a power law increase with an exponent that becomes larger as $R/\dot{\gamma}$ increases. For large $R/\dot{\gamma}$, there are avalanches whose linear size ($\sim S^{1/2}$) becomes larger than L_z . This means that avalanches will be macroscopically observable even in the large system size limit ($L_z \rightarrow \infty$).

(c) The stress distribution across the system is qualitatively different for small or large $R/\dot{\gamma}$. At small values of this parameter, stress fluctuations decrease rapidly in the sample as system thickness increases. Yet for large $R/\dot{\gamma}$, stress maintains

long-wavelength spatial fluctuations that decay more slowly with L_z .

(d) An examination of the flow curve of the system reveals clear signs of velocity weakening when $R/\dot{\gamma} \neq 0$. This velocity wakening behavior is seen to be less detectable as the system size (namely, L_z) increases.

These results make clear the differences between two different yielding behavior of materials, one that can be qualified as smooth, or ductile, and a second one that is fragile, and akin to a system that localizes deformation and can produce earthquakes at the largest scales.

II. THE MODEL

Our approach is based on the modeling presented in [9]. The applied deformation on the system corresponds to an affine displacement field **u** of the form $\mathbf{u} \sim (\gamma z, 0, 0)$, which gives the (xz) component of the strain as the only one that is nonzero. The average strain is, therefore, γ and is the quantity that is externally controlled. At the mesoscopic level, the system responds with a local strain $e_{xz}(\mathbf{r})$ that, for simplicity, will be referred to as $e(\mathbf{r})$, and which is the main variable in the problem. Note that $\overline{e(\mathbf{r})} = \gamma$, where the overline indicates a spatial average. It has to be remarked that although we are using a "scalar" model (in the sense that it is defined by a single-component deformation field), full three-dimensional phenomenology is encoded in the form of the elastic interaction, as explained below. For a full three-dimensional approach to the problem, see [31].

The temporal evolution equation for $e(\mathbf{r})$ is assumed to be of the overdamped form, therefore giving $\dot{e}(\mathbf{r})$ as proportional to a generalized force acting on $e(\mathbf{r})$. This force has two contributions: a local part f_{loc} and an elastic interaction part f_{int} . The elastic interaction term describes the effect that a change in $e(\mathbf{r})$ has on a different point \mathbf{r}' . It can be expressed as an integral over the whole system of the form

$$f_{\rm int}(\mathbf{r}) = \int d\mathbf{r}' G(\mathbf{r}, \mathbf{r}') e(\mathbf{r}').$$
(1)

The elastic kernel $G(\mathbf{r}, \mathbf{r}')$ depends on the symmetry of the applied deformation, and it also includes the effect of other elastic deformations with different symmetry that are not explicitly taken into account in the formalism. In the case in which the deformation is a pure shear, *G* takes the form that is known as the Eshelby interaction, which has a quadrupolar symmetry. However, for the present simple shear deformation case, it takes a slightly different form [32], namely,

$$G_{\mathbf{q}} = \frac{q_x^2 + q_y^2}{q_x^2 + q_y^2 + q_z^2}.$$
 (2)

Note that this represents a long-range interaction with an asymptotic form in real space, $\sim 1/r^3$, as the standard Eshelby interaction, yet the spatial symmetry is dipolar in this case.

The local part of the force, f_{loc} , models the disordered nature of the system, and it consists of a potential with many minima at different values of *e*, representing different locally stable equilibrium configurations of our amorphous system. The forces f_{loc} are taken totally uncorrelated at different points in the system, namely, there are no correlations in the relative positions of the energy minima at different points. The form of



FIG. 2. Sketch of the local potential for the variable e, at some position in the sample. The width Δ of each potential well is a stochastic variable uniformly distributed between 0.2 and 2.2. Note that different spatial positions in the system have different, uncorrelated forms of this function.

the local potentials, $V_{\text{loc}}(e)$, determining the forces, $f_{\text{loc}}(e) \equiv -dV_{\text{loc}}/de$, is effectively chosen in the following way. At each spatial position, we take a local equilibrium value of e, namely, e_0 , in such a way that $V_{\text{loc}}(e) = (e - e_0)^2/2$, and the force $f_{\text{loc}}(e) = (e - e_0)$. When, during the dynamical evolution, $|e - e_0|$ becomes greater than some value Δ , the local potential well is supposed to destabilize and the local minimum e_0 shifts to a new value. This produces a typical local potential $V_{\text{loc}}(e)$, as sketched in Fig. 2. Once the form of the local and interaction forces is defined, the model is completely defined and ready to be simulated.

Although our model focuses on the dynamics of strain and stress, it is probably useful to add a comment with respect to the elastic energy in the system. When the applied strain increases, and as long as there are no avalanches, strain energy increases and is stored in the local potential energy of each site, V_{loc} . On the other hand, when an avalanche occurs, some elastic energy is dissipated (via the overdamped dynamics). In a realistic situation of a shear stress material, this dissipation would produce some heating on the sample, or even the radiation of elastic waves in a seismicity context.

The exact form of f_{loc} , and its relative amplitude with respect to the elastic term f_{int} , have an effect, for instance, on the precise value of σ_c . However, we have observed that the kind of effects we are interested in do not depend qualitatively on details in the exact definition of these forces. Therefore, we are quite convinced that the results we are obtaining are robust with respect to changes of the system parameters.

The system is driven by externally enforcing the average value of e to be equal to the applied strain γ . It has to be stressed that in the present case, no additional parameter is included to enforce the driving condition. This is contrary to the standard situation in depinning models where a spring with adjustable stiffness is added to drive the system. In the simulations shown below, γ is increased quasistatically, so each avalanche in the system occurs at a fixed value of strain. Along the simulations, the stress σ in the system is calculated as the average value of the local forces, $\sigma \simeq \overline{f_{\text{loc}}}$ (note that the system average value of f_{int} vanishes). When an avalanche occurs, σ drops by some amount $\Delta \sigma$, which is taken as a measure of the avalanche size S.

We consider the possibility of relaxation in the system, which is the property that will drive strain localization and "earthquakelike" properties of the avalanches. Relaxation represents a tendency in the system to achieve more relaxed states as time progresses, and will compete with the driving, which pushes the system away from its most relaxed states. It has been introduced in a variety of manners in different models of yielding. Here we use the implementation originally presented in [9], which was able to generate a realistic phenomenology of shear bands, when driving the system at a finite $\dot{\gamma}$. Relaxation is introduced through a parameter *R*, which represents a backaction of the local forces on the form of the local potentials trying to uniformize the stress across the system. This is implemented by the following evolution scheme [9]:

$$\dot{e}_0 = R\nabla^2(e_0 - e). \tag{3}$$

The value of *R* has to be considered in relation to the strain rate $\dot{\gamma}$. Although we work in the limit of very low $\dot{\gamma}$, in such a way that strain can be considered constant during the evolution of any single avalanche, there is a strong dependence of the system behavior on the ratio $R/\dot{\gamma}$, which will be the crucial parameter to describe the dynamics of the model. The only other free parameter is the system size, which we choose to be *L* along the *x* and *y* directions, and L_z along *z*.

Before going to the results that we find, we want to mention the status of the proposed kind of modeling with respect to other approaches used for similar problems. The kind of modeling that is presented here was initiated by Bulatov and Argon [33] and then applied to other kinds of problems (see, for instance, Refs. [34,35]). Its application to the problem of yield stress materials with relaxation was initiated in [9]. It was not clear in the beginning whether this kind of simulations provided the same results as other more standard approaches at that time, particularly the so-called elastoplastic models (EPMs) [3]. The question was settled with a detailed comparison of both kinds of modeling [36], showing that our approach provides the same qualitative results as those of EPMs. In particular, some unexpected dependencies of the results on the form of local potentials V_{loc} (particularly, whether they are smooth or have a "cuspy form") were also related to the properties of EPMs when studied with stress-dependent yielding rates [36]. Therefore, although the present approach is not the only possibility for a mesoscopic model, it is one that is compatible with other well-accepted modeling techniques, in particular EPMs.

The use of the relaxation parameter R and its physical meaning deserve further discussion. The effect of R [through Eq. (3)] is to allow the potential energy landscape (coded in the values of e_0 at different positions) to adapt to the strain distribution in the system (determined by e), reaching progressively lower energy (i.e., more stable) configurations. Although the present implementation is very specific, the general description of a "relaxation" term is quite general and has been implemented in a number of ways in other kinds of modeling (see discussion and references in [3], Sec. V C). We have shown in a number of papers (particularly [29,37]) that this relaxation (which can also be interpreted as a viscoelastic term in the equations of motion) is able to reproduce a great deal of the phenomenology of earthquakes when applied to two-dimensional systems. Actually, one of the points of the present paper is to see how, in three-dimensional samples, the



FIG. 3. (a) Avalanche probability distribution P(S) as a function of avalanche size S for four different values of L_z as indicated (the values of L used along x and y are also reported, but they are irrelevant because, in all cases, they are large enough so as to avoid size effects due to L). (b) The same data as in (a), but plotted as $P(S)S^r$ vs S/L^{ζ} , with $\tau = 1.5$, $\zeta = 1.2$.

relaxation can produce the localization of the deformation in a two-dimensional "fault."

III. RESULTS

We first describe the results without relaxation, i.e., the case $R/\dot{\gamma} = 0$. They are consistent with those obtained with slightly different numerical models (particularly those called elastoplastic models), although typically in those cases a cubic geometry has been used, instead of the slab geometry we are using here.

Along the simulation, the avalanches are characterized by the stress drop they produce in the system. This defines the size S of each avalanche. From a sequence of S values, we define the avalanche size distribution P(S) as (proportional to) the number of avalanches of sizes between S and S + dS. In Fig. 3(a), we show the avalanche size distribution in samples with different values of L_z . As explained before, the value of L along the x and y directions is large enough in such a way that the results are independent of the precise value of L. Yet, for completeness, the values of L are also reported in Fig. 3. As the scaling in Fig. 3(b) shows, the results can be fitted by a power law with a large size cutoff, in the form

$$P(S) \sim S^{-\tau} \exp(-S/S_{\max}), \tag{4}$$

with $\tau \simeq 1.5$, and $S_{\text{max}} \sim L_{z}^{\sim 1.2}$. The value $\tau \simeq 1.5$ is the one expected for long-range elastic interaction, which in fact is active here due to the structure of the elastic kernel *G*.



FIG. 4. Upper and lateral view of the system indicating the sites that were affected by a few successive large avalanches (each avalanche is represented by a different color). Individual avalanches are more or less localized along the *z* direction; however, successive avalanches appear at random *z* positions (system size is $128 \times 128 \times 32$). In a very long run, avalanches uniformly cover the full three-dimensional system.

The observation of the sites that participate in each individual avalanche (Fig. 4) gives further insight into their structure. Active sites for individual avalanches tend to be spread along the x and y directions, but remain rather localized along z. The sites participating in an avalanche do not form a connected set, but typically consist of a number of disconnected pieces. This is an effect that is enhanced by the long-range elastic interactions. When observing a temporal sequence of avalanches, it is seen that they do not occur around the same definite value of z, but, rather, they appear randomly in the system. Therefore, when considering the deformation induced by many avalanches, this is uniform across the system, with no sign of strain localization.

Although the largest avalanches in the system diverge in size as $L_z \rightarrow \infty$, the *x*-*y* span actually becomes progressively smaller compared with L_z as L_z increases. In fact, we observed that the number of positions in the *x*-*y* plane affected by an avalanche behaves similarly to the avalanche size itself [38]. Then the span along the *x*-*y* plane can be estimated as $\sim S_{\text{max}}^{1/2}$, which increases sublinearly with L_z (since $S_{\text{max}} \sim L_z^{1.2}$). This implies that avalanches become "vanishingly small" compared with system thickness as this is increased, and in the large thickness limit, the yielding of the system will be uniform and smooth. We will come back to this fact when comparing with the results for $R/\dot{\gamma} > 0$.

Now we present the results obtained by including relaxation in the model. In this case, the system is characterized by a competition between the rate of local relaxation and the strain rate. Namely, the control parameter will be $R/\dot{\gamma}$. In Fig. 5, we plot the sites affected by a few consecutive



FIG. 5. Same as Fig. 4 for finite relaxation $(R/\dot{\gamma} = 0.3)$. The main observed difference is the localization of the deformation around a "fault" that is formed near a definite *z* value (system size is $128 \times 128 \times 16$). In a very long run, avalanches uniformly cover the *x*, *y* projection of the system, but they remain localized around the fault in the *Z* direction.

avalanches with a size larger than some minimum. One remarkable feature that is observed at finite relaxation compared to the unrelaxed case is that deformation is no longer uniform in the system, but localizes in a nearly two-dimensional "fault" at some particular value of z. This is qualitatively observed in the localization of the avalanches as observed in the lateral view in Fig. 5, and it is also more quantitatively established in Fig. 6, where we plot the accumulated deformation that has affected each individual layer of the system within a long run. While this deformation is seen to be uniform when R = 0, it is clearly localized in z when $R/\dot{\gamma} \neq 0$.

This localization effect is reminiscent of the shear band localization observed in cases of finite $\dot{\gamma}$. However, in the



FIG. 6. Normalized accumulated deformation $\gamma(z)$ at each z value during a long run, as a function of the relaxation parameter $R/\dot{\gamma}$. The deformation is seen to be uniform when R = 0, but it becomes localized when $R/\dot{\gamma} \neq 0$.



FIG. 7. Avalanche probability distribution P(S) as a function of avalanche size *S* for three different values of L_z and different values of $R/\dot{\gamma}$ as indicated (curves corresponding to different values of $R/\dot{\gamma}$ were vertically displaced, for better comparison).

present case, the argument of a "Maxwell construction" does not apply, as we do not have uniform sliding within the shear band. Yet the reason to have localization is rather similar: the region of the system in which avalanches occur maintains a state of low relaxation, which produces a local slower critical stress, whereas the rest of the system which is fully blocked had plenty of time to relax and therefore it has a large critical stress. As a consequence, deformation continues to occur systematically at the same spatial positions. In this respect, it has to be mentioned that this localization process is highly hysteretic when the value of $R/\dot{\gamma}$ is changed. The curves shown in Fig. 6 were obtained starting at $R/\dot{\gamma} = 0$ and progressively increasing it. If, from the final large value of $R/\dot{\gamma}$, this is reduced back to zero, it is very difficult to return to a situation of uniform deformation because blocked parts of the sample already have a large critical stress and are not ready to flow even if relaxation is suppressed. The recovery of a uniform deformation occurs via a very slow process that was quantitatively described in [39-41] in the context of shear bands at finite strain rate.

We now analyze how the avalanche size distribution is affected by relaxation. In Fig. 7, we show size distributions for different values of L_z and $R/\dot{\gamma}$ (here again, L_x and L_y are taken to be large enough, so as not to affect the results). By comparing curves with the same value of L_z , we see that in the presence of relaxation, the power law distribution is maintained, roughly with the same value of τ , but the increase of $R/\dot{\gamma}$ makes the avalanche distribution broader, with a value of S_{max} that increases with $R/\dot{\gamma}$. In Fig. 8, the avalanche distributions are plotted as a function of S/L_z^{ζ} , where the values of ζ are chosen to collapse the curves corresponding to the same $R/\dot{\gamma}$, namely, we have an *R*-dependent ζ value. This dependence is shown in the inset.

We see that ζ shows a strong increase as a function of $R/\dot{\gamma}$. It starts at $\zeta \simeq 1.2$ for R = 0, reaching and even surpassing the value $\zeta = 2$ when $R/\dot{\gamma} \sim 1$. The value $\zeta = 2$ is somewhat of a critical value, in the following sense. Since the linear size of an avalanche in the *x*-*y* plane can be estimated as $\sim S^{1/2}$, we see that if $\zeta < 2$, the maximum linear size of the avalanches becomes negligible compared to the system thickness L_z as L_z increases. On the contrary, if $\zeta > 2$, this maximum size



FIG. 8. The data of the previous figure scaled according to S/L_z^{ζ} . The values of ζ used to scale the data for different values of $R/\dot{\gamma}$ are indicated in the inset. The line is a guide to the eye.

stays larger than L_z in the thermodynamic limit. We conclude that there is a critical value of $R/\dot{\gamma}$ below which avalanches become vanishingly small when $L_z \rightarrow \infty$, whereas above that value, there are avalanches that have a linear extent in the x-y plane that is much larger than L_z . We suggest this is an indication (in the large system size limit) of a transition between smooth, ductile yielding and a fragile behavior in which the effect of individual avalanches continues to be observable even if L_z is increased arbitrarily. This is one of the main findings of the present work.

IV. ROUGHNESS OF THE SPATIAL DISTRIBUTION OF STRESS

In the previous section, we showed that avalanches become vanishingly small in the thermodynamic limit $(L_z \rightarrow \infty)$ in the absence of, or for low values of, relaxation, whereas there are "macroscopic" avalanches (linear size larger than L_z) for large enough relaxation. We will explore the consequence of this fact in the stress distribution across the sample.

In the simulation, we have access to the local value of stress at every time, which is given by $\sigma(\mathbf{r}, t) = e(\mathbf{r}, t) - e_0(\mathbf{r}, t)$. To be able to extract some manageable information from this, we do the following. First, we work with an average of this quantity over the *z* coordinate. This generates a function [that we continue to call $\sigma(x, y, t)$ for simplicity] that is much simpler to analyze and also eliminates the large differences that would certainly appear between points close to or away from the yielding plane (at least for cases where spatial localization occurs). Therefore, we take the function $\sigma(x, y, t)$ at fixed times as an indication of the spatial fluctuation of stress.

As before, we first consider the R = 0 case. In Fig. 9, we see the function $\sigma(x, y)$ at some fixed time for systems with different values of L_z . An examination of these figures suggests that the fluctuation of σ across the system is reduced as L_z is increased, which is compatible with that stated in the previous section. In order to have a more quantitative confirmation of this behavior, we calculate the structure factor of these configurations and plot the results in Fig. 10. We see a rather flat structure, indicating the lack of correlations of σ among different parts of the sample. Moreover, the typical value of the structure factor decreases as L_z increases,



FIG. 9. Instantaneous stress fluctuation across the sample, for different values of $L_z = 8$, 16, 32 ($L_x = L_y = 128$), in the absence of relaxation.

indicating that the stress distribution become progressively more uniform in thicker samples. We find a decay of the typical fluctuation of stress w across the sample that follows $w \sim L_z^{-1}$.



FIG. 10. Structure factor S(q) of the configurations in Fig. 9 ("q = 1" corresponds nominally to the maximum q-value allowed by the discrete mesh). The structure factor is flat and its value decreases as L_z^{-1} .



FIG. 11. Same as Fig. 10 for $L_z = 4$, 8, 16 (we depict a 128 × 128 piece of a system with $L_x = L_y = 512$), for $R/\dot{\gamma} = 1$.

Now we do the same analysis in a case with $R/\dot{\gamma} = 1$. We see the distribution of σ across the system in Fig. 11. It is apparent that in the present case, there are noticeable correlations between the values of σ at different spatial positions. To be more quantitative about this point, we again resort to the analysis of the structure factor. The curves corresponding to the three values of L_z (obtained by averaging a large number of configurations to minimize the statistical error) are shown in Fig. 12. For large q values, the structure factor decreases with L_z , as before. However, for the lowest values of q, a remarkable independence of the structure factor on L_z is observed. This independence occurs in a range of small q that corresponds to a spatial region of the order of L_z , and it is related to the correlation that is visible in the snapshots in Fig. 11. We should stress, however, that the independence of S(q) with L_z observed at low q does not imply that the stress distribution has a finite width in the limit $L_z \rightarrow \infty$. In fact, our model does not sustain persistent long-range correlations as the relaxation mechanism is a diffusive process that tends to uniformize the stress in the system. In other words, if we go to a limit in which $R/\dot{\gamma} \rightarrow \infty$ (for instance, setting $\dot{\gamma} = 0$ while R remains finite), the system configuration will evolve



FIG. 12. Structure factor of the configurations in Fig. 11 (an average over many equivalent configurations was performed). Now the structure factor reaches an L_z -independent value as $q \rightarrow 0$.

towards a state of uniform stress. Yet the correlations in stress that are observed originate in the existence of large (linear size $\sim L_z$) avalanches in the presence of relaxation. In the next section, we correlate this effect with the existence of "velocity weakening" in the system.

V. FLOW CURVE AND VELOCITY WEAKENING

The spatial correlations of stress that we observe for finite $R/\dot{\gamma}$ are a consequence of the velocity weakening at the fundamental level caused by relaxation. Velocity weakening is the underlying mechanism behind a reentrant flow curve as that in Fig. 1(b); yet, the weakening part of this curve (dotted line) is not observed in the thermodynamic limit, in the same way that a reentrant liquid-gas isotherm of a system displaying a liquid-gas transition is not observed, since it is screened by the coexistence of phases. However, in finite systems, the velocity weakening behavior can be observed and this is what we want to show here.

Our simulations are run in a quasistatic limit in which avalanches are instantaneous events. This means that we are exploring the $\dot{\gamma} \rightarrow 0$ limit of the full flow curve of the system. Yet, we still have, as an independent parameter, the ratio $R/\dot{\gamma}$. To discuss the flow curve, it is more instructive to think of R as fixed and $\dot{\gamma}$ as variable. By collecting results from our simulations, we constructed the plot in Fig. 13. The velocity weakening effect in this curve is clearly visible. Yet the effect becomes less visible as L_{z} is increased. This is not surprising. A negatively sloped velocity-stress dependence is a mechanically unstable situation that at finite strain rates, is screened by the breakdown of spatial homogeneity. In fact, the localization of deformation along the z direction that was described in the previous section is a manifestation of this effect, where the unstable behavior is cured through a "Maxwell construction mechanism" in which most of the system is blocked, and only a very thin region yields. But even when the deformation is localized in a single layer (which plays the role of a "seismic fault"), the yielding may have a velocity weakening character that the system will screen by producing strong inhomogeneities in the x-y plane. This is actually the reason for the strong fluctuations that we have observed in the previous section for the stress across the system.



FIG. 13. Flow curve of our quasistatic simulations showing the velocity weakening effect. Results for $L_z = 4$ (red) and $L_z = 16$ (blue). Note how the velocity weakening effect becomes less pronounced for the thicker sample. Also, because our simulations are quasistatic, we do not observe the increase of σ with $\dot{\gamma}$ that typically occurs at finite $\dot{\gamma}$, as sketched in Fig. 1.

This result reinforces the idea of two different yielding regimes in the thermodynamic limit, depending on the extent of relaxation. One is the smooth regime, occurring for low or no relaxation, in which the effect of individual avalanches is washed out in the thermodynamic limit, and the stress distribution becomes asymptotically uniform. The second regime, occurring at large relaxation, displays some avalanches that are larger in linear size that the system thickness L_z , no matter how large this is, and also spatial fluctuations in the values of stress that are intimately related to the velocity weakening nature of the yielding at short scales.

VI. SUMMARY AND CONCLUSIONS

In this work, we have presented numerical simulations of a model for a yield stress material under quasistatic deformation. The aim was to investigate under which circumstances the deformation is smooth and uniform at large scales, or jerky and localized. The main result is that the answer to this question depends on the amount of "relaxation" that is included in the model. This relaxation also affects other characteristics of the model, such as the global flow curve.

For no or low levels of relaxation, the avalanches that are responsible for the plastic deformation of the system scale with system size in such a way that they become comparatively small as the system size increases. Also, they appear all across the system and the observed deformation is spatially uniform and smooth in the large system size limit. In this case, the flow curve of the material is monotonous, and the overall behavior of the material can be described as "ductile."

However, when large levels of relaxation are present, the deformation localizes in a very thin layer, and the maximum size of the observed avalanches is comparable to the system thickness no matter how large this is, meaning that individual avalanches may have macroscopic effects even for very large system sizes. This phenomenology strongly resembles the situation occurring in geological faulting. The very thin layer when deformation localizes is just the seismic fault, and avalanches with an effect that propagates all across the system are the earthquakes. All this phenomenology is associated to an underlying velocity weakening behavior of the system, which is responsible for both the localization of the deformation in a thin fault and the fact that avalanches within this fault have noticeable effects at the system size scale.

ACKNOWLEDGMENT

The author acknowledges discussions with Alberto Rosso and Giuseppe Petrillo on topics related to the content of this work.

- P. Coussot, Yield stress fluid flows: A review of experimental data, J. Non-Newtonian Fluid Mech. 211, 31 (2014).
- [2] D. Bonn, M. M. Denn, L. Berthier, T. Divoux, and S. Manneville, Yield stress materials in soft condensed matter, Rev. Mod. Phys. 89, 035005 (2017).
- [3] A. Nicolas, E. E. Ferrero, K. Martens, and J.-L. Barrat, Deformation and flow of amorphous solids: Insights from elastoplastic models, Rev. Mod. Phys. 90, 045006 (2018).
- [4] D. S. Fisher, Collective transport in random media, Phys. Rep. 301, 113 (1998).
- [5] M. Kardar, Nonequilibrium dynamics of interfaces and lines, Phys. Rep. 301, 85 (1998).
- [6] G. Picard, A. Ajdari, L. Bocquet, and F. Lequeux, Simple model for heterogeneous flows of yield stress fluids, Phys. Rev. E 66, 051501 (2002).
- [7] P. D. Olmsted, Perspectives on shear banding in complex fluids, Rheol. Acta 47, 283 (2008).
- [8] T. Divoux, M. A. Fardin, S. Manneville, and S. Lerouge, Shear banding of complex fluids, Annu. Rev. Fluid Mech. 48, 81 (2016).
- [9] E. A. Jagla, Strain localization driven by structural relaxation in sheared amorphous solids, Phys. Rev. E 76, 046119 (2007).
- [10] J. Lin, E. Lerner, A. Rosso, and M. Wyart, Scaling description of the yielding transition in soft amorphous solids at zero temperature, Proc. Natl. Acad. Sci. USA 111, 14382 (2014).
- [11] S. Karmakar, E. Lerner, and I. Procaccia, Statistical physics of the yielding transition in amorphous solids, Phys. Rev. E 82, 055103(R) (2010).
- [12] K. A. Dahmen, Y. Ben-Zion, and J. T. Uhl, Micromechanical Model for Deformation in Solids with Universal Predictions for Stress-Strain Curves and Slip Avalanches, Phys. Rev. Lett. 102, 175501 (2009).
- [13] B. Tyukodi, S. Patinet, S. Roux, and D. Vandembroucq, From depinning transition to plastic yielding of amorphous media: A soft-modes perspective, Phys. Rev. E 93, 063005 (2016).
- [14] The meaning of fragile here is unusual. Usually, fragile means that a sample will break suddenly when deformation is increased. In the present case, the sample does not break, but individual avalanches of plastic rearrangements produce observable effects even in samples of arbitrarily large size.
- [15] C. H. Scholz, *The Mechanics of Earthquakes and Faulting* (Cambridge University Press, Cambridge, 2002).
- [16] A. Kabla and G. Debregeas, Local Stress Relaxation and Shear Banding in a Dry Foam under Shear, Phys. Rev. Lett. 90, 258303 (2003).
- [17] H. Rehage and H. Hoffman, Viscoelastic surfactant solutions: model systems for rheological research, Mol. Phys. 74, 933 (1991).
- [18] J.-F. Berret, D. C. Roux, and G. Porte, Isotropic-to-nematic transition in wormlike micelles under shear, J. Phys. II France 4, 1261 (1994).

- [19] V. Schmitt, F. Lequeux, A. Pousse, and D. Roux, Flow Behavior and Shear Induced Transition near an Isotropic/Nematic Transition in Equilibrium Polymers, Langmuir 10, 955 (1994).
- [20] C. Grand, J. Arrault, and M. E. Cates, Slow transients and metastability in wormlike micelle rheology, J. Phys. II France 7, 1071 (1997).
- [21] M. M. Denn, Issues in viscoelastic fluid mechanics, Annu. Rev. Fluid Mech. 22, 13 (1990).
- [22] C. Tang, H. Peng, Y. Chen, and M. Ferry, Formation and dilatation of shear bands in a Cu-Zr metallic glass: A free volume perspective, J. Appl. Phys. **120**, 235101 (2016).
- [23] A. L. Greer, Y. Q. Cheng, and E. Ma, Shear bands in metallic glasses, Mater. Sci. Eng.: R: Rep. 74, 71 (2013).
- [24] S. Ogata, F. Shimizu, J. Li, M. Wakeda, and Y. Shibutani, Atomistic simulation of shear localization in CuZr bulk metallic glass, Intermetallics 14, 1033 (2006).
- [25] C. Zhong, H. Zhang, Q. P. Cao, X. D. Wang, D. X. Zhang, U. Ramamurty, and J. Z. Jiang, Deformation behavior of metallic glasses with shear band like atomic structure: A molecular dynamics study, Sci. Rep. 6, 30935 (2016).
- [26] R. Burridge and L. Knopoff, Model and theoretical seismicity, Bull. Seismol. Soc. Am. 57, 341 (1967).
- [27] J. M. Carlson, J. S. Langer, and B. E. Shaw, Dynamics of earthquake faults, Rev. Mod. Phys. 66, 657 (1994).
- [28] Z. Olami, Hans Jacob S. Feder, and K. Christensen, Self-Organized Criticality in a Continuous, Nonconservative Cellular Automaton Modeling Earthquakes, Phys. Rev. Lett. 68, 1244 (1992).
- [29] E. A. Jagla, F. P. Landes, and A. Rosso, Viscoelastic Effects in Avalanche Dynamics: A Key to Earthquake Statistics, Phys. Rev. Lett. **112**, 174301 (2014).
- [30] Note that we have stated that we work in the limit $\dot{\gamma} \rightarrow 0$. The fact that $R/\dot{\gamma}$ is finite means that relaxation is also a very slow process in the system. In other words, we can say that during the duration of any avalanche, *R* and $\dot{\gamma}$ do not have any effect on the system.
- [31] E. A. Jagla, Tensorial description of the plasticity of amorphous composites, Phys. Rev. E **101**, 043004 (2020).
- [32] For a derivation of this kernel, see the Appendix of I. Fernández Aguirre and E. A. Jagla, Phys. Rev. E 98, 013002 (2018).
- [33] V. V. Bulatov and A. S. Argon, A stochastic model for continuum elasto-plastic behavior. I. Numerical approach and strain localization, Modell. Simul. Mater. Sci. Eng. 2, 167 (1994).
- [34] T. Lookman, S. R. Shenoy, K. O. Rasmussen, A. Saxena, and A. R. Bishop, Ferroelastic dynamics and strain compatibility, Phys. Rev. B 67, 024114 (2003).
- [35] S. Kartha, J. A. Krumhansl, J. P. Sethna, and L. K. Wickham, Disorder-driven pretransitional tweed pattern in martensitic transformations, Phys. Rev. B 52, 803 (1995).
- [36] E. E. Ferrero and E. A. Jagla, Criticality in elastoplastic models of amorphous solids with stress-dependent yielding rates, Soft Matter 15, 9041 (2019).

- [37] E. A. Jagla and A. B. Kolton, A mechanism for spatial and temporal earthquake clustering, J. Geophys. Res. 115, B05312 (2010).
- [38] This occurs because, typically, no more than a single site with the same *x*-*y* coordinates participates in an avalanche, and additionally because every participating site typically moves to the next potential well of the local potential, and no more than that.
- [39] E. A. Jagla, Shear band dynamics from a mesoscopic modeling of plasticity, J. Stat. Mech. (2010) P12025.
- [40] Y. Shi, M. B. Katz, H. Li, and M. L. Falk, Evaluation of the Disorder Temperature and Free-Volume Formalisms via Simulations of Shear Banding in Amorphous Solids, Phys. Rev. Lett. 98, 185505 (2007).
- [41] D. D. Alix-Williams and M. L. Falk, Shear band broadening in simulated glasses, Phys. Rev. E 98, 053002 (2018).