by Meek and Saxe that, during the passage of the streamer tip in long sparks, there is a cusp-shaped luminosity that extends out radially at least 10-cm beyond the streamer axis as sensed by their photomultiplier. After this passes, the streamer channel is largely dark until the return stroke illuminates a channel of some 3-mm radius or less. In the case of lighting discharge, the radius of photographed vigorous steps with heavy currents runs out to some 10 meters. In such strokes, the channel at the cloud end of the stroke may remain visible for some little time. In this case the pilot leader invisible to the camera advances 20 to 200 meters with an active conducting channel of  $R=10$  cm illuminated by the return stroke from ground some 10 milliseconds later, and has as well the expanded channel of positive space charge of radius  $R'$ which, with its roughly  $5\times10^{10}$  electrons per cm<sup>3</sup>, is illuminated only by the step flash from the cloud end within some 10—100 microseconds. After ionization and illumination, the step ionization of radius  $R'$  decays, and it is only the original channel of  $R=10$  cm which has sufficient conductivity owing to the field  $X<sub>c</sub>$  to carry the return stroke.

It will be noted from Table III that the observed values of what might be  $R'$  are less than those computed. This is not surprising, for, as radial expansion continues, the time rate of ionization and accompanying excitation decline, and what is preceived by photomultiplier or photographic plate corresponds to values considerably less than  $R'$  depending on the sensitivity of the detector. Amin could observe only at a value of radius greater than 0.05 cm from the streamer axis. That he did not observe luminosity between 0.05 and 0.1 cm is not surprising. That, however, there was a transient luminosity after the intensely luminous peak passed the slit, is shown by the shoulder of luminosity following the tip and of such shape as only to be accounted for by ionization and excitation occurring long-after even an especially broad tip had passed the slit. The resolving power of Meek and Saxe's system was not such as to have revealed the details of the fine structure of the tip luminosity as observed by Amin, while their radial transient tip expansion could well be observed with their heavy currents. From the evidence presented, even though observed luminosity and calculated values  $R$  agree only in order of magnitude, it is believed that the nature of the transient tip shape and luminosity is accounted for. The transient space charge expansion described is illustrated schematically in Fig. 2, which is self-explanatory.

PHYSICAL REVIEW VOLUME 94, NUMBER 2 APRIL 15, 1954

# The Scattering of Electromagnetic Waves by Turbulent Atmospheric Fluctuations\*

F. VILLARS AND V. F. WEISSKOPF

Department of Physics and Lincoln Laboratory, Massachusetts Institute of Technology, Cambridge, Massachusetts (Received December 7, 1953)

The statistical theory of turbulence is applied to the problem of density fluctuations in the troposphere and the ionosphere. For suitable wavelengths, for which the so-called similarity region (Kolmogoroff spectrum) of the spectrum of turbulence is relevant, a closed formula can be given for the scattering cross section. It contains as only parameter the turbulent power dissipation S, and its angular dependence is given by  $(\sin \frac{1}{2}\theta)^{13/3}$ ,  $\theta$  being the scattering angle. The values of S required to explain ionospheric scattering are in excellent agreement with values found from investigations of meteor trails. Tropospheric data cannot be fitted with the assumptions of dry-air turbulence alone. The inference is that humidity fluctuations play an essential part in tropospheric scattering. A preliminary study of these latter fluctuations gives satisfactory results. Further investigations (and experimental data) are needed, however, to work out a quantitative theory.

IGHT waves are scattered by random fluctuations  $\overline{\mathbf{L}}$  of the refractive index. In what follows we derive the scattering of elementary waves by random fiuctuations which are produced by turbulent perturbations. The general idea underlying this study has been suggested by a number of authors, especially Megaw' and Booker.<sup>2</sup> It will be shown that, under certain conditions, the scattering produced by these fluctuations

can be expressed in terms of only one parameter, the turbulent energy S dissipated per cm' per sec. The conditions of validity of this relation are well fulfilled for the scattering of meter waves in the  $E$  layer of the ionosphere; for tropospheric scattering other parameters such as the inhomogeneity of potential temperature and specific humidity play an important part.

#### A. SIMPLE DERIVATION OF THE SCATTERING FORMULA

#### l. Scattering Cross Section

In this section we derive the expressions by simple qualitative arguments and leave the exact derivations for Sec. B.

<sup>\*</sup>The research in this document was supported in part by the U. S. Army, Navy, and Air Force under contract with the Massachusetts Institute of Technology.<br><sup>1</sup> E. C. S. Megaw, Nature **166**, 1100 (1950), and Proc. Inst.

Elec. Engrs. (London) 100, 7 (1953).<br><sup>2</sup> H. Booker and W. E. Gordon, Proc. Inst. Radio Engrs. 38,

<sup>401</sup> (1950).

We calculate the scattering which a light beam with the wave vector  $\mathbf{k}_0$  suffers in a region of space of volume V, which contains a medium whose dielectric constant  $\epsilon$  fluctuates:  $\epsilon = \epsilon_0 + \Delta \epsilon$ , where  $\Delta \epsilon$  is a function of space. We assume that the time variations are sufficiently slow and do not affect the scattering. Then the scattered electric field amplitude  $(E=E_s e^{i(\omega t - k_1 \cdot r)})$  at a distance R from the scattering volume  $(R \gg V^{\frac{1}{3}})$  is given by

$$
E_s = \frac{E_0}{4\pi R\lambda^2} \left| \int_V d\mathbf{r} \Delta \epsilon(\mathbf{r}) e^{i\mathbf{K} \cdot \mathbf{r}} \right| \sin\chi. \tag{1}
$$

Here  $E_0$  is the electric-field amplitude of the incident wave  $(E_{\text{inc}}=E_0e^{i(\omega t - k_0 \cdot r)})$ ,  $\lambda$  is the wavelength divided by  $2\pi$ ,  $\chi$  is the angle between the direction of the incident field and the scattering, and

$$
K = |\mathbf{K}| = |\mathbf{k}_0 - \mathbf{k}_1| = 2k \sin \frac{1}{2}\theta, \tag{2}
$$

where

$$
k = |\mathbf{k}_0| = |\mathbf{k}_1| = \lambda^{-1}.
$$

We introduce the cross section  $\sigma d\Omega$  for the scattering per unit volume of the scattering volume into the solid angle  $d\Omega$ :

and get

$$
\sigma d\Omega\!=\!(R^2/V)\,(E_s{}^2\!/E_0{}^2),
$$

$$
\sigma d\Omega = |M|^2 \sin^2 \chi d\Omega / V \lambda^4 (4\pi)^2, \tag{3}
$$

with

and

$$
M = \int d\mathbf{r} \Delta \epsilon(\mathbf{r}) e^{i\mathbf{K} \cdot \mathbf{r}}.
$$
 (3a)

For the ionosphere, we assume that  $\Delta \epsilon$  is due to fluctuations in electron concentration, proportional to the density fluctuations  $\Delta \rho$  of the carrier medium:

$$
\Delta \epsilon = g \Delta \rho / \rho, \tag{4}
$$

$$
g = -\omega_N^2/\omega^2; \quad \omega_N^2 = (1/\epsilon_0)(Ne^2/m), \tag{5}
$$

where  $N$  is the density of free electrons in the medium. In the troposphere, we will have to cope with fluctuations in density, potential temperature and humidity. Because of the large dipole moment of the water molecule,<sup>3</sup> this latter contribution is particularly important. The dielectric constant may be written as

$$
\epsilon - 1 = 1.55 \times 10^{-4} \frac{p}{T} \left( \frac{\text{mb}}{\text{o}_{\text{K}}} \right) + 0.74 \frac{p'}{T^2} \left( \frac{\text{mb}}{\text{o}_{\text{K}}^2} \right). \tag{6}
$$

In (6),  $\dot{p}$  is the pressure of air,  $p'$  the partial pressure of water vapor. Since  $p/T \sim \rho$ , we best express  $\Delta \epsilon$  in terms of density fluctuations:

terms of density fluctuations:  
\n
$$
\Delta \epsilon = 0.45 \Delta \rho \left(\frac{g}{cm^3}\right) + 3.4 \times 10^8 \left[\frac{\Delta \rho'}{T} - \frac{\rho'}{T^2} \Delta T\right] \left(\frac{g/cm^3}{\sigma_K}\right).
$$
\n
$$
\text{on} \quad \text{on} \quad (6a)
$$

In (6a),  $\rho$  refers to the density of air,  $\rho'$  to the density of water vapor, both in  $g/cm<sup>3</sup>$ . In what follows, we shall

<sup>8</sup> C. M. Crain, Phys. Rev. 74, 691 (1948).

develop a description of the density fluctuations  $\Delta \rho$ from a statistical dynamics of turbulence. Since the problems involved in the discussion of the humidity fluctuations [second term in  $(6a)$ ] are of rather different nature we shall discuss them separately, and presently we shall discuss the density fluctuations in dry and isentropic air only.

#### 2. Density Fluctuations in Homogeneous Turbulence

First we connect the density fluctuations with the velocity fluctuations by means of Bernoulli's law: Local velocity differences  $\Delta v$  will be accompanied by pressure differences  $\Delta p$ , and density differences  $\Delta p$ , according to

$$
\Delta \rho / \rho \cong \Delta p / p = (\Delta v)^2 / v_M^2, \tag{7}
$$

where  $v_M^2$  is the average square of the molecular velocity. We now concentrate our attention on the velocity fluctuations under conditions of homogeneous turbulence. ' We may picture this situation as follows: there are external causes which constantly produce large eddies of a certain dimension  $L_0$  and velocity  $v_0$ . These eddies divide soon into smaller eddies, say, of size  $L_1 = \alpha L_0$ , with  $\alpha < 1$ . The velocity of these smaller eddies embedded in the larger ones is  $v_1$ . (Measurement of  $v_1$ is relative to the environment, i.e. , relative to the motion of the large eddy.) These eddies divide again into smaller ones  $L_2 = \alpha^2 L_0$ , with a velocity  $v_2$  relative to their surrounding. In this dividing process, the energy fed into the largest eddy is transferred to smaller and smaller ones. A constant amount of energy per volume and time is fed into the larger eddies from an outside source. The dividing process ends with that eddy size at which the effect of the molecular viscosity is large enough to dissipate the energy into heat.

We now establish the quantitative relations which exist in this process. The energy per unit volume contained in one eddy is of the order  $\rho v_n^2$ . The lifetime of such an eddy will be of the order  $t_n = L_n/v_n$ . Hence, the eddies of size  $L_n$  lose energy to smaller ones (size  $L_{n+1}$ ) with a rate (energy/time volume):

$$
S_n \sim \rho v_n^3 / L_n. \tag{8}
$$

Since we have a stationary process, the loss  $S_{n-1}$  of the eddies one size larger to those of size  $n$  must be equal to  $S_n$ ; hence,  $S_n$  is a constant independent of n:

$$
\rho v_n^3 / L_n = S. \tag{8a}
$$

S is the energy transmitted through the eddies from the energy source (larger eddies) all the way down to the smaller ones. The order of magnitude of  $S$  is given by the transfer from the largest eddies:

$$
S \cong \rho v_0^3 / L_0. \tag{8b}
$$

<sup>4</sup> The theory of turbulence as used here was irst suggested by A. N. Kolmogoroff [Compt. rend. acad. sci. U.R.S.S. 30, 301 (1941)]. Our treatment is closest to the ideas of G. F. v. Weizsäcker [Z. Physik 124, 614 (1948)] and W. Heisenberg [Z. Physik 124, 628 (1948)].

We can use the magnitude of  $S$  in order to determine the size of the smallest eddies: the energy per cm' and sec dissipated by molecular viscosity is given by  $\sim_{\eta} (dv/dx)^2$ , where  $dv/dx$  is the rate of change of v per unit length. The smallest eddy is the one in which this energy reaches S. Since  $(dv/dx) \sim (v_n/L_n)$ , we get for the smallest eddy, whose velocity is  $v<sub>S</sub>$  and whose L is  $L_s$ ,

$$
\eta v_S^2/L_S^2 \cong S.
$$

Using the relation  $(8a): \rho v_s^3/L_s = S$ , and also (8b), we get

$$
L_0/L_S = (\rho v_0 L_0/\eta)^{3/4}.
$$
 (9)

In words, the ratio of the sizes of the largest to the smallest eddies is the  $\frac{3}{4}$ -power of the Reynolds number, associated with the eddy-producing large scale kinematics.

We can identify the velocities  $v_n$  with the magnitude  $\Delta v$  used in (6). We consider regions in space of the linear size  $L, L$  fulfilling the relation

$$
L_S \le L \le L_0. \tag{10}
$$

Let us then assume that  $L_n = L$ ; then  $\Delta v = v_n$  is the average deviation from its surrounding of the velocity in that region. We then get from (6) the average density deviation  $(\Delta \rho)_L$  from its surrounding in a region L:

$$
(\Delta \rho)_L = \rho (v_n/v_M)^2 = \rho (v_0/v_M)^2 (L/L_0)^{\frac{2}{3}}.
$$
 (11)

Let us illustrate these results by a few examples:

(a) In the froposphere, at 10-km height, we have  $\rho = 3.6 \times 10^{-4}$  g cm<sup>-3</sup>,  $\eta = 2 \times 10^{-4}$  g cm<sup>-1</sup> sec<sup>-1</sup>. We estimate the order of magnitude of size and velocities in turbulent gusts to 1 km and 10 m/sec, respectively. Hence we may put

$$
v_0 = 10^3 \text{ cm/sec}, \quad L_0 = 10^5 \text{ cm}, \tag{12}
$$

and, therefore,

$$
R=1.8\times10^8
$$
,  $L_0/L_s=1.55\times10^6$ ,  $L_s=0.65$  mm,  
 $S=0.36$  erg cm<sup>-3</sup> sec<sup>-1</sup>.

The smallest eddies are a fraction of a cm large.

(b) In the  $E$ -layer of the *ionosphere* at 100-km height, we have  $\rho = 2 \times 10^{-9}$  g cm<sup>-3</sup>. The viscosity is independent of pressure, and, therefore, again  $\eta = 2 \times 10^{-4}$ g cm<sup>-1</sup>  $sec^{-1}$ . From the very scant data about velocities and eddy sizes we infer roughly:

$$
v_0 = 5 \times 10^3
$$
 cm sec<sup>-1</sup>,  $L_0 = 5 \times 10^5$  cm, (13a)

and hence'

$$
R = 2.5 \times 10^4, \quad L_0/L_S = 2 \times 10^3, \quad L_S = 2.5 \times 10^2 \text{ cm},
$$
  
 
$$
S = 0.5 \times 10^{-3} \text{ erg cm}^{-3} \text{ sec}^{-1}. \tag{13b}
$$

We now proceed to discuss the *humidity fluctuations*. A rough estimate of the order of magnitudes indicates

that the effects of humidity in the troposphere are probably much stronger than the effects of the fluctuations of the density of the dry air. As can be inferred from (7) and (19), the average mean square fluctuation  $\langle \Delta \rho^2 \rangle_{\rm Av}$  of the latter density is of order  $\rho (v_0 / v_M)^4$ . With the data as given by (12), and with  $v_M \sim 4 \times 10^4$  cm/sec, we get

$$
(1/\rho^2)\langle \Delta \rho^2 \rangle_{\text{Av}} \sim 0.4 \times 10^{-6}.\tag{14}
$$

We infer from (6a) that for  $\rho' \sim 10^{-2} \rho$ ,  $T \sim 300^{\circ}$ , a  $\langle \Delta \rho'^2 \rangle_{\rm Av}/\rho'^2$  or  $\langle \Delta T^2 \rangle_{\rm Av}/T^2$  of magnitude (14) produces a  $\langle \Delta \epsilon^2 \rangle_{\text{Av}}$  twice as large as the dry-air density fluctuation of (14). Meteorological evidence indicates, however, that the fluctuations  $\langle \Delta \rho'^2 \rangle_{\text{Av}}$  and  $\langle \Delta T^2 \rangle_{\text{Av}}$  may be considerably larger: Measured  $\langle \Delta T^2 \rangle_{\text{AV}}/T^2$  values are of order of magnitude  $10^{-5}$ ;<sup>6</sup> unfortunately humidity measurements are scarce, but indicate<sup>7</sup> fluctuations that may be considerably larger. Practically nothing is known about the process of turbulent mixing of inhomogeneously humid air. Thus no information about short-range correlation of humidity in turbulent air is available. With some knowledge of the dynamics of turbulence it should, however, be possible to calculate approximate correlation functions from measured large-scale fluctuations (over dimensions of the order  $L_0$ ). Thus, for instance, we may assume that, in an atmosphere with a negative gradient of potential temperature, convection produces large "wet" eddies (they need not be actual clouds) embedded in comparatively dryer air. The dissipation of this "surplus humidity" into the surrounding medium is then to follow closely the dissipation of the kinetic energy of the wet eddy by the breaking-up process described above. This leads to the tentative conclusion of a spectrum of humidity fluctuations very similar to the spectrum of the fluctuations of  $v^2$ , and this again is equal to the density fluctuation spectrum. Should this indeed be the case—a close study of this point is under way—the knowledge of large-scale humidity variations  $(\Delta \rho')_0$  would serve to determine the humidity fluctuations  $(\Delta \rho')_L$  in a volume of order L compared to its surrounding.

In fact we would get from (11)

$$
(\Delta \rho')_L / (\Delta \rho')_0 = (\Delta \rho)_L / (\Delta \rho)_0 = v_n^2 / v_0^2 = (L / L_0)^{\frac{3}{2}}.
$$
 (14a)

Unfortunately very little is known about the large-scale humidity fluctuations  $(\Delta \rho')_0$ .

# 3. Calculation of the Scattering Cross Section

We now determine the magnitude  $M$  appearing in (1) and defined by (2). With  $(4)<sup>8</sup>$  we write

$$
M = \frac{g}{\rho_0} \int d\mathbf{r} \Delta \rho(\mathbf{r}) e^{i\mathbf{K} \cdot \mathbf{r}},\tag{14b}
$$

<sup>&</sup>lt;sup>5</sup> This estimate is in good agreement with the values given by C. deJager [Mem. soc. roy. sci., Liège, Series 4, XXI, 223 (1952)]<br>and obtained from an analysis of meteor trails.

<sup>&</sup>lt;sup>6</sup> See Electrical Engineering Research Laboratory, University<br>of Texas, Reports Nos. 47, 1950, and 53, 1953 (unpublished).<br>T Electrical Engineering Research Laboratory, University of<br>Texas, Reports Nos. 54, 1951, and 6-01

<sup>&</sup>lt;sup>8</sup> In case of tropospheric scattering, the dry-air term of (6a) gives  $g=0.45\rho(g/cm^3)$ .

where  $\rho_0$  is the average density, and we try to determine this magnitude on the basis of our picture of homogeneous turbulence. These concepts will be applicable only if the length  $K^{-1}$  lies between  $L_S$  and  $L_0$ :

$$
L_S < (2k \sin \frac{1}{2}\theta)^{-1} < L_0.
$$
 (15)

The integral appearing in (14) is a Fourier coefficient of the density fluctuations. We can expand the density  $\rho(r)$  in a Fourier series within the scattering volume V (which we may assume to be a cube):

$$
\rho(\mathbf{r}) = \frac{1}{V} \sum_{i} \rho(\mathbf{k}_{i}) e^{-i\mathbf{k}_{i} \cdot \mathbf{r}},
$$
\n(16)

with the inversion

$$
\rho(\mathbf{k}_{i}) = \int_{V} d\mathbf{r} \rho(\mathbf{r}) e^{i\mathbf{k}_{i} \cdot \mathbf{r}}.
$$
 (17)

It is then clear that  $M$ , as given by (14), is nothing but

$$
M = (g/\rho_0)\rho(\mathbf{K}).\tag{18}
$$

We shall now have to establish the relation between  $\rho(\mathbf{k})$  and the quantity  $(\Delta \rho)_L$  introduced in Sec. (2). Now  $(\Delta \rho)_L$  is built up by the fluctuations whose spatial period is of order L. Because of the random nature of these fluctuations, different wavelengths will not interfere, and we have

$$
(\Delta \rho)_{L}^{2} = \frac{1}{V^{2}} \sum_{k \sim 1/L} | \rho(\mathbf{k}) |^{2} \cong \frac{1}{V^{2}} \left(\frac{V}{2\pi^{2}}\right) \int_{k'}^{k''} k^{2} dk | \rho(\mathbf{k}) |^{2}.
$$

The interval  $k''-k'$  is of order  $1/L$  itself and centered 1. Cross Section about  $K=1/L$ .  $(V/2\pi^2)k^2dk$  is the number of Fourier components with wavelength between  $k$  and  $k+dk$ . This gives us roughly:

 $(\Delta\rho)_L{}^2\!\!\cong\!\!\frac{1}{6\pi^2}\frac{1}{VL^3}|\,\rho\left(1/L\right)|{}^2$ 

Hence,

$$
|M|^2 = 6\pi^2 \left(\frac{g}{\rho_0}\right)^2 V L^3 (\Delta \rho)_{L^2} \Bigg|_{L=1/K},
$$

and inserting  $(\Delta \rho)_L^2$  from (11), we get

$$
|M|^2 = 6\pi^2 g^2 V \left(\frac{v_0}{v_M}\right)^4 \frac{L^{13/3}}{L_0^{4/3}} \bigg|_{L=1/K}.
$$

Inserting this into (3) gives

$$
\sigma d\Omega \cong g^2 \left(\frac{v_0}{v_M}\right)^4 \frac{L^{13/3}}{L_0^{4/3}} \frac{\sin^2 \chi}{\chi^4} d\Omega, \tag{20}
$$

where a factor of order unity  $(3/8)$  is omitted. Setting we get  $L = K^{-1}$  and using (2) gives, finally,

$$
\sigma = \frac{g^2}{L_0} \left(\frac{v_0}{v_M}\right)^4 \left(\frac{\tilde{\lambda}}{L_0}\right)^{1/3} (2 \sin{\frac{1}{2}\theta})^{-13/3} \sin^2\!\chi. \tag{21}
$$

The validity of this expression is limited by the condition (10).It is only applicable as long as

$$
L_0^{-1} < 2k \sin\frac{1}{2}\theta < L_S^{-1}.\tag{22}
$$

We find that the turbulent conditions enter the cross section only in the form  $v_0^4/L_0^{4/3}$ ; hence, the cross section is proportional to  $(S)^{4/3}$  and depends only upon the energy dissipation S.

We note the characteristic angular dependence  $(\sin \frac{1}{2}\theta)^{-13/3}$ . The dependence upon the wavelength is different in the troposphere and in the ionosphere. In the former,  $e^2$  is independent of  $\lambda$ , and we get  $\sigma \sim \lambda^{1/3}$ . In the latter, we get from (Sb) a wavelength dependence  $\sim \lambda^{13/3}$ .

If the conditions (22) are no longer fulfilled, we expect deviations from (21). The case where  $L = (2k \sin{\frac{1}{2}\theta})^{-1}$ is comparable to or smaller than  $L_S$  is of special interest. In this case, the eddies of that size are attenuated by molecular dissipation, and  $(\Delta \rho)_L$  is then smaller than (11). In particular, it will fall off stronger than  $L^{2/3}$ with decreasing  $L$ . Let us put instead of  $(11)$ , for example,

$$
(\Delta \rho)_L = \rho (v_0/v_M)^2 (L/L_0)^{2/3} F(L/L_S), \tag{23}
$$

where  $F(x)$  is a function which is equal to unity for  $x \gg 1$  and which falls off rapidly for  $x < 1$ . We then get a factor  $|F(\lambda/2L_s \sin \frac{1}{2}\theta)|^2$  multiplied into the cross section  $(21)$ . Hence, the scattering falls off more strongly with decreasing  $\lambda$  or increasing  $\theta$  if  $(2k \sin^2 \theta)^{-1}$  $\langle L_s.$ 

# B. OUANTITATIVE THEORY

With Eqs. (1) and (2), we can write the cross section (1) as

$$
\sigma d\Omega = \frac{g^2}{(4\pi)^2} \frac{1}{\lambda^4 \rho^2 V} \left| \int_V d\mathbf{r} \Delta \rho(\mathbf{r}) e^{i\mathbf{K} \cdot \mathbf{r}} \right|^2 \sin^2 \chi d\Omega.
$$

Now  
\n
$$
\frac{1}{\rho^2} \Big| \int_V d\mathbf{r} \Delta \rho(\mathbf{r}) e^{i\mathbf{K} \cdot \mathbf{r}} \Big|^2 = \int d\mathbf{r} e^{i\mathbf{K} \cdot \mathbf{r}} \frac{1}{\rho^2} \int d\mathbf{R} \Delta \rho(\mathbf{R}) \Delta \rho(\mathbf{R} + \mathbf{r})
$$
\n
$$
= \int d\mathbf{r} e^{i\mathbf{K} \cdot \mathbf{r}} V \frac{\langle \Delta \rho^2 \rangle_{\text{Av}}}{\rho^2} C(\mathbf{r}). \quad (24)
$$

Defining the Fourier-transformed  $C(\mathbf{k})$  of the correlation function  $C(\mathbf{r})$  by

$$
C(\mathbf{r}) = \left(\frac{1}{2\pi}\right)^3 \int d\mathbf{k} e^{-i\mathbf{k} \cdot \mathbf{r}} C(\mathbf{k}),\tag{25}
$$

 $(19)$ 

$$
\sigma d\Omega = \frac{g^2}{(4\pi)^2} \frac{1}{\lambda^4} \left( \frac{\langle \Delta \rho^2 \rangle_{\text{Av}}}{\rho^2} \right) C(\mathbf{K}) \sin^2 \chi d\Omega, \tag{26}
$$

with  $K=2k \sin \frac{1}{2}\theta$ .

## 2. Velocity and Density Spectrum in the Statistical Theory of Turbulence'

We shall use a Fourier series expansion of the velocity field as given in Eqs. (16) and  $(17)$ ;<sup>10</sup> this enables us to write the fundamental Navier-Stokes equation in the form

$$
\dot{\mathbf{v}}_{k} + i \sum_{\mathbf{k}'} (\mathbf{k} \cdot \mathbf{v}_{k'}) \mathbf{v}_{k-k'} + \frac{\eta}{\rho} k^2 \mathbf{v}_{k} + i \frac{\mathbf{k}}{\rho} p_k = 0.
$$
 (27)

We shall assume an incompressible medium

$$
(\mathbf{k} \cdot \mathbf{v}_{\mathbf{k}}) = 0. \tag{27a}
$$

This may sound paradoxical in view of our attempt to calculate the density fluctuation  $\Delta \rho$ . Actually our assumption merely implies that the divergence free part of the velocity field is dynamically much more important than the irrotational components. It also amounts to assuming that the potential energy stored in density variations is much smaller than the kinetic energy of the vortices. With this in mind, we shall now simply calculate the pressure fluctuations  $\Delta \phi$ , and then calculate  $\Delta \rho$  with the aid of

$$
p(\mathbf{r}) = p_0[\rho(\mathbf{r})/\rho_0]^\gamma, \tag{28}
$$

 $\rho_0$ ,  $\rho_0$  being average values of density and pressure, and  $\gamma = C_p/C_v = 1.4$  (for diatomic molecules). Multiplying (27) with **k** and realizing that for  $k \neq 0$ 

 $p_k = \Delta p_k$ 

we have

$$
\Delta p_{\mathbf{k}}\!=\!-(\rho_0/k^2)\!\sum_{\mathbf{k}'}(\mathbf{k}\!\cdot\!\mathbf{v}_{\mathbf{k}'})\!\left(\mathbf{k}\!\cdot\!\mathbf{v}_{\mathbf{k}\!-\!\mathbf{k}'}\right)
$$

and hence

$$
\Delta \rho_k = -(\rho_0/\gamma \rho_0)^2 k^{-2} \sum_{\mathbf{k}'} (\mathbf{k} \cdot \mathbf{v}_{\mathbf{k}'})(\mathbf{k} \cdot \mathbf{v}_{\mathbf{k} - \mathbf{k}'})
$$
 (29)

So we now turn our attention to the velocity field  $v(k)$ . Equation (27) gives, upon multiplication with  $v(-k)$ ,

$$
-\frac{\rho}{2}\frac{d}{dt}(\mathbf{v}_{\mathbf{k}}\cdot\mathbf{v}_{-\mathbf{k}})=i\rho\sum_{k'}(\mathbf{k}\cdot\mathbf{v}_{\mathbf{k'}})(\mathbf{v}_{-\mathbf{k}}\cdot\mathbf{v}_{\mathbf{k}-\mathbf{k'}})+\eta k^2(\mathbf{v}_{\mathbf{k}}\cdot\mathbf{v}_{-\mathbf{k}}).
$$
 (30)

We recognize  $\frac{1}{2}\rho v(\mathbf{k}) \cdot v(-\mathbf{k})$  as the mean kinetic energy density carried by the wave number k, since

$$
\frac{1}{2}\rho\langle\mathbf{v}^2(\mathbf{x})\rangle_{\mathrm{Av}} = \frac{1}{2}\rho \sum_k (\mathbf{v}_k \cdot \mathbf{v}_{-k}).
$$

Summing over all  $k$ 's in (30), we get the total *power* balance:

$$
S = -\frac{d}{dt} \left( \frac{\rho}{2} \sum_{k} (\mathbf{v}_{k} \cdot \mathbf{v}_{-k}) \right)
$$
  
=  $i \rho \sum_{k, k'} (\mathbf{k} \cdot \mathbf{v}_{k'}) (\mathbf{v}_{-k} \cdot \mathbf{v}_{k-k'}) + \eta \sum_{k} k^{2} (\mathbf{v}_{k} \cdot \mathbf{v}_{-k}).$  (30a)

paper (reference 4).<br><sup>10</sup> We also use  $v_k = v(k)$ .

The first term on the right-hand side describes the *transfer* of kinetic energy to different wave numbers  $\mathbf{k}'$ , the second term the *viscous dissipation* of energy. We shall now define a stationary situation within a region  $(L_0^3)$  completely enclosed in the domain over which (27) holds: this region then contains a largest eddy of linear dimension  $\overline{L_0}$  and characteristic velocity  $v_0$ . This eddy is coupled to the outside of that region by virtue of the "transfer" term in (30a). We assume that the effect of this coupling is to maintain the average velocity  $v_0$  of this eddy by means of a power supply  $S_0$ . Provided  $S_0$  is constant in time, an equilibrium will be reached in which the average  $v(k)$  is also independent of time.

From (30a) we see that  $S_0$  must be of the form

$$
S_0 = \alpha \rho (v_0^3 / L_0),
$$

with  $\alpha$  an absolute constant (independent of the values of  $v_0$  and  $L_0$ ), depending only on the large scale geometry of the power-supplying region outside  $L_0^3$ . Subsequently, we shall put  $\alpha=1$  and *define*  $S_0$  by

$$
S_0 = \rho(v_0^3/L_0). \tag{31}
$$

If the Reynolds number  $R_0 = \rho v_0 L_0 / \eta$  is sufficiently large, the dissipative term will enter into play only for wave numbers  $k\gg1/L_0$ . It will affect eddies of size  $L_s$ and velocity  $v_s$ , for which the "local" Reynolds number has dropped to I:

$$
\rho L_s v_s / \eta = 1. \tag{32}
$$

Once wave numbers of order  $1/L_s$  are reached, the frictional power dissipation replaces the power transfer to higher wave numbers. So we get another equation characterizing  $v_s$  and  $L_s$ :

$$
S_0 = \eta(v_S/L_S)^2. \tag{32a}
$$

We shall use  $(32)$  and  $(32a)$  to *define* the two quantities  $v_S$  and  $L_S$ :

$$
v_S = (S_0 \eta / \rho^2)^{1/4}; \quad L_S = (\eta^3 / S_0 \rho^2)^{1/4}.
$$
 (33)

We see that the three "external" parameters  $S_0$ ,  $\eta$ ,  $\rho$ set an absolute scale of length and velocities for the problem of energy transfer and dissipation in turbulent flow.

To carry this idea through in a more quantitative way, let us define a *spectral intensity* distribution  $F(k)$ for the velocities:

$$
\frac{1}{2} \sum_{\mathbf{k}} \mathbf{v}_{\mathbf{k}} \cdot \mathbf{v}_{-\mathbf{k}} = \int_0^\infty dk F(k). \tag{34}
$$

(In this definition, isotropy of the velocity spectrum has been assumed.) The dissipative term in (30a) is then

$$
2\eta \int_0^\infty dk k^2 F(k). \tag{35}
$$

<sup>&#</sup>x27;Much of the content of this section is to be found in G. K. Batchelor's recent publication on the theory of homogeneous turbulence [*The Theory of Homogeneous Turbulence* (Cambridge<br>University Press, London, 1953)] as well as Heisenberg's origina

Since (35) represents the total power absorbed, we from dimensional arguments: have, with  $y=kL_s$ :

$$
S_0 = (2\eta/L_S^3) \int_0^\infty dy y^2 F(L_S^{-1}y).
$$

$$
1 = \int_0^\infty dy \, 2y^2 E(y). \tag{36}
$$

# Determination of  $E(y)$  or  $F(k)$

# (a) Inertial range

Provided  $L_s \ll L_0$ , there is a region  $k \ll 1/L_s$ , in which the shape of the velocity spectrum  $F(k)$  is determined by the transfer mechanism and not affected by the viscous dissipation. In this range  $F(k)$  should be independent of  $\eta$ , or  $v_s^2 L_s E(kL_s)$  independent of  $\eta$ . Clearly this can only be satisfied with a power law for  $E$ , since  $L<sub>g</sub>$  contains  $\eta$ . Putting

$$
E(kL_s) = \text{const} \times (kL_s)^n,
$$

we get, with  $(33)$ :

$$
k^n v_S^2 L_S^{n+1} \propto k^n \eta^{\frac{1}{2} + \frac{3}{4}(n+1)},
$$

and independence of  $\eta$  gives  $n = -5/3$ .

To normalize  $F$  approximately, we can tentatively assume that this spectrum holds for all  $k$  values between  $1/L_0$  and  $1/L_s$  and that beyond both these limits the spectrum is cut off. On the high wave number side, the viscosity is certainly to act as a cutoff, whereas on the low end the power source produces naturally eddies of a certain maximum size  $\overline{L_0}$ . With these assumptions the normalization condition (36) gives us:

$$
const \times (2\eta/S_0) \int_{1/L_0}^{1/L_S} dk k^{\frac{1}{3}} = 1.
$$

(We see here that the exact shape of the spectrum at the lower end is immaterial as far as the normalization goes.) Hence,

and

const
$$
\approx \frac{2}{3}(S_0/\eta)L_S^{4/3} = \frac{2}{3}(S_0/\rho)^{2/3},
$$
  
 $F(k) = \frac{2}{3}(S_0/\rho)^{\frac{2}{3}}k^{-5/3}.$  (37)

# (b) Tail of Velocity Spectrum

For this region no really good solution of the problem has been given. The simplest way to deal with the problem is to consider the transfer of kinetic energy from small to large wave numbers as a damping effect on the motion of the large eddies. Thus the concept of an *eddy viscosity*  $\eta'(k)$  is introduced to describe the drain of energy on an eddy of size  $k^{-1}$  by eddies of size  $\langle k^{-1}$ . If we assume that  $\eta'(k)$  can at all be expressed in terms of  $F(k')$   $(k' < k)$ , a unique expression results

$$
S_0 = (2\eta/L_S^3) \int_0^\infty dy y^2 F(L_S^{-1}y).
$$
  $\eta'(k) = \text{const} \times \rho \int_k^\infty dk' [F(k')/k'^3]^{\frac{1}{2}}.$  (38)

Using (32a), we see that  $F(L_S^{-1}y)=v_S^2L_SE(y)$  defines The value of the constant can be estimated from the a *universal* function  $E(y)$  normalized to 1:  $v(k)$  a applies of the argument that  $\eta'(k)$  should be large compared argument that  $\eta'(k)$  should be large compared with  $\eta$ for  $k<1/L_s$  and becoming equal  $\eta$  at  $k=1/L_s$ . Hence,

(36) 
$$
\eta = \text{const} \times \rho \int_{1/L_S}^{\infty} dk' [F(k')/k'^3]^{\frac{1}{2}};
$$

with (37), this gives:

with

and

$$
const = (4/3)\sqrt{\frac{3}{2}} \cong 1.
$$

We can now picture our stationary power transfer as a dissipation process whereby the energy  $S_k$  dissipated in eddies  $k' \leq k$  or transferred to the velocity components  $v_{k'}$ , with  $k' > k$  is given by

$$
S_k = 2\left[\eta'(k) + \eta\right] \int_0^k dk' k'^2 F(k'). \tag{40}
$$

The condition of stationary transfer implies that no energy is accumulated in a given wave number interval and hence that  $S_k$  is *independent of* k and equal to  $S_0$ . With (38) we thus get the following equation for  $F(k)$ :

$$
S_0 = 2 \left\{ \rho \int_k^{\infty} dk' [F(k')/k'^3]^{\frac{1}{2}} + \eta \right\} \int_0^k dk'' k''^2 F(k''). \quad (41)
$$

The solution of this equation is found to  $be<sup>11</sup>$ 

$$
F(k) = (8S_0/9\rho)^{2/3}k^{-5/3}(1+(k/k_S)^4)^{-4/3},\qquad(42)
$$

$$
k_S = (3S_0 \rho^2/8\eta^3)^{1/4}.
$$

As we see, for  $k \ll k_s$ , we find again our result (37) with <sup>a</sup> slight—and for our purpose, immaterial —change of the normalization constant. Also the value  $k<sub>S</sub>$  is essentially identical with  $1/L_s$  as given in (33). For  $k\gg k_s$ ,  $F(k)$  behaves like  $k^{-7}$ :

$$
F(k) \cong 8^{-2/3} (S_0 \rho/\eta)^2 k^{-7}.
$$
 (43)

It might be interesting to compare here  $k_s$  with  $k_0 = 1/L_0$ and  $k_M = 1/L_M$ ,  $L_M$  being the *mean free path* of the molecules of the gas: From simple gas kinetics, we have

$$
(\eta/\rho)=(2\pi/3)v_M/k_M,
$$

$$
R = v_0 L_0 / v_M L_M
$$
 (Reynolds number).

From (33) we then obtain:

$$
k_S = R^{4/3} k_0, \tag{44a}
$$

[see also (9)] and 
$$
k_M = (S_M/S_0)k_S,
$$
 (44b)

 $11$  The solution of Eq. (41) is due to J. Bass, Compt. rend. 228, 22 (1949).

where we have introduced  $S_M = \rho v_M^3 k_M$  in analogy to  $S_0$ .

Equation (44a) shows that  $k_s \gg k_0$  if only the Reynolds number is large enough.

Equation (44b) decides the more subtle question whether hydrodynamic concepts apply at all in the discussion of the dissipation process.  $k_M \gg k_S$  is a necessary condition for it, and it holds only provided the power supply  $S_0$  is not too large. We shall discuss the numerical aspect of these conditions in Sec. C.

Our analysis of the velocity spectrum is now essentially completed, and we shall presently return to a discussion of the density fluctuations as given by Eq. (29).

The correlation function  $C(r)$  introduced in (24) may be expressed in terms of the  $\rho_k$  ( $\rho_k = \Delta \rho_k$  for  $k \neq 0$ ):

$$
\langle \Delta \rho^2 \rangle_{\text{Av}} C(\mathbf{r}) = \langle \Delta \rho (\mathbf{R}) \Delta \rho (\mathbf{R} + \mathbf{r}) \rangle_{\text{Av}}
$$
  
= 
$$
\frac{1}{V} \int_{V} d\mathbf{R} \Delta \rho (\mathbf{R}) \Delta \rho (\mathbf{R} + \mathbf{r})
$$
  
= 
$$
\sum_{k} \langle \rho_{k} \rho_{-k} \rangle_{\text{Av}} e^{i\mathbf{k} \cdot \mathbf{r}}
$$

[Note: The remaining averages —once the space averages are carried out—are time (phase) averages. ) With (29) and  $\langle \rho^2 \rangle_{\text{Av}} \cong \rho_0^2$ , we get

$$
C(\mathbf{r}) = \left(\frac{\rho_0}{\gamma p_0}\right)^2 \sum_{\mathbf{k}} k^{-4} e^{i\mathbf{k}\cdot\mathbf{r}} \sum_{\mathbf{k'}\mathbf{k'}} \langle (\mathbf{k}\cdot\mathbf{v}_{k'}) (\mathbf{k}\cdot\mathbf{v}_{k-k'})
$$

$$
\times (\mathbf{k}\cdot\mathbf{v}_{k''})(\mathbf{k}\cdot\mathbf{v}_{k'k-k}) \rangle_{\text{Av}}.
$$

The sum  $\sum_{k'k''}$  can be collapsed if we notice that only the combinations

(a) 
$$
k'=k''
$$
,  $k-k'=k-k''$ ,  
(b)  $k'=k-k''$ ,  $k-k'=k''$ 

give nonzero contributions in the average. Thus  $\sum_{\mathbf{k}'\mathbf{k}''}$ reduces to

$$
2 \sum_{k'} \langle (\mathbf{k} \cdot \mathbf{v}_{k'}) \mathbf{(k} \cdot \mathbf{v}_{-k'}) \rangle_{\text{Av}} \times \langle (\mathbf{k} \cdot \mathbf{v}_{k-k'}) \mathbf{(k} \cdot \mathbf{v}_{k'-k}) \rangle_{\text{Av}}.
$$

Furthermore, since the incompressibility implies  $(\mathbf{k} \cdot \mathbf{v}_k)$  $=0$ , we have:

$$
\langle (\mathbf{k} \cdot \mathbf{v}_{k'}) \mathbf{(k} \cdot \mathbf{v}_{-k'}) \rangle_{\text{av}} = \langle (\mathbf{k}_{\perp} \cdot \mathbf{v}_{k'}) \mathbf{(k}_{\perp} \cdot \mathbf{v}_{-k'}) \rangle_{\text{av}} = \frac{1}{2} | \mathbf{v}_{k'} |^2 k_{\perp}^2,
$$

where

$$
\mathbf{k}_{\perp} = \mathbf{k} - \mathbf{n}'(\mathbf{k} \cdot \mathbf{n}'), \quad \mathbf{n}' = \mathbf{k}'/k',
$$

and, therefore,

$$
k_{\perp}^2 = \left[k^2k'^2 - (\mathbf{k} \cdot \mathbf{k}')^2\right]/k'^2.
$$

This gives

$$
\begin{aligned}\n &\langle \langle \Delta \rho^2 \rangle_{\text{Av}} / \rho^2 \rangle C(\mathbf{r}) \\
 &= 2 \left( \rho_0 / \gamma \rho_0 \right)^2 \sum_k k^{-4} e^{i\mathbf{k} \cdot \mathbf{r}} \sum_{\mathbf{k'}} \langle |\mathbf{v}_{k'}|^2 \rangle_{\text{Av}} \langle |\mathbf{v}_{k-k'}|^2 \rangle_{\text{Av}} \\
 &\quad \times \frac{\left[ k^2 k'^2 - (\mathbf{k} \cdot \mathbf{k'}) \right]^2}{4k^2 (\mathbf{k} - \mathbf{k'})^2}\n \end{aligned}
$$

Since we need only the Fourier-transformed  $C(\mathbf{k})$ of  $C(r)$ , [see (25) and (26)], we write the **k** sums in terms of Fourier integrals with the aid of the substitution

$$
\sum{}_{\bf k}\rightarrow V(2\pi)^{-3}\int d{\bf k}.
$$

Introducing at the same time  $F(k)$  by (34a), we get

$$
\begin{aligned}\n&\langle \langle \Delta \rho^2 \rangle_{\text{Av}} / \rho^2 \rangle C(\mathbf{r}) \\
&= \pi (\rho_0 / \gamma \rho_0)^2 (2\pi)^{-3} \int d\mathbf{k} k^{-4} e^{i\mathbf{k} \cdot \mathbf{r}} \\
&\times \int d\mathbf{k}' F(k') F(|\mathbf{k} - \mathbf{k}'|) \bigg( \frac{k^2 k'^2 - (\mathbf{k} \cdot \mathbf{k}')^2}{k^2 (\mathbf{k} - \mathbf{k}')^2} \bigg)^2,\n\end{aligned}
$$

or, in view of (26):

$$
\langle \Delta \rho^2 \rangle_{\text{Av}} / \rho^2 C(\mathbf{k}) = \pi (\rho_0 / \gamma \rho_0)^2 k^{-4} \int d\mathbf{k}' F(k')
$$

$$
\times F(|\mathbf{k} - \mathbf{k}'|) \left( \frac{k^2 k'^2 - (\mathbf{k} \cdot \mathbf{k}')^2}{k^2 (\mathbf{k} - \mathbf{k}')^2} \right)^2. \quad (45)
$$

We shall first evaluate (45) for a case

$$
k_0 \ll k \ll k_S. \tag{46}
$$

In that case, we shall use the spectrum  $F(k)$  as given by (37) and use a straight cutoff at  $k \ge k_s$ , as well as for  $k \leq k_0$ . We shall see, however, that with the assumptions of  $(46)$ , the k' integral is essentially independent of the two limits. The k' integral in (45)—let us call it  $J(k)$  is most easily evaluated in bipolar coordinates:

$$
k' = kr_1, \quad |{\bf k} - {\bf k}'| = kr_2, \quad d{\bf k}' = 2\pi k^3 r_1 r_2 dr_1 dr_2.
$$

This gives for  $J(k)$ :

$$
J(k) = (8S_0/9\rho)^{4/3} (2\pi k^{-1/3}) \int \int dr_1 dr_2 (r_1 r_2)^{-14/3}
$$

$$
\times \left[ (r_1^2 - r_2^2)^2 - 2(r_1^2 + r_2^2) + 1 \right]^2 / 16
$$

$$
= (8S_0/9\rho)^{4/3} (2\pi k^{-1/3}) H(k_0/k; k_s/k). \tag{47}
$$

 $H$  is a dimensionless integral and with (46) approximately an absolute constant. So we have

$$
(\langle \Delta \rho^2 \rangle_{\text{Av}} / \rho^2) C(\mathbf{k}) \cong 2\pi^2 (\rho_0 / \gamma \rho_0)^2 (S_0 / \rho)^{4/3} k^{-13/3} H. \quad (48)
$$

An approximate integration gives for  $H$ , assuming (46),

$$
H = 1.2 + 1.6 \frac{k}{2k_s}^{7/3} + \cdots. \tag{49}
$$

(In all practical cases,  $k$  is probably very much larger than  $k_0$  and not so far from  $k_s$ . We, therefore, did not evaluate corrections  $\propto k_0/k$  in H.)

In the limit of very high wave numbers  $k$ ,

$$
k \gg k_S, \tag{50}
$$

238

the main contribution to  $J(k)$  will arise from the shaded regions of Fig. 1. So we may write

$$
J(k) = \int d\mathbf{k}' F(k') F(|\mathbf{k} - \mathbf{k}'|) \sin^4\alpha
$$
  

$$
\approx 2 \int_{k' < ks} d\mathbf{k}' F(k') F(k) \sin^4\alpha
$$

Introducing again dimensionless variables  $z: k' = kz$ , we get

$$
J(k) \cong 2(8S_0/g\rho)^{4/3}k^{-1/3}(k_S/k)^{16/3}\int_{z\leq k_S/k}dz z^{-5/3}\sin^4\alpha,
$$

with

$$
k_S \gg k \quad (50), \quad \sin^2 \alpha \cong \sin^2 \varphi \text{ (see Fig.1)},
$$

and hence

$$
\int dz z^{-5/3} \sin^4 \alpha \approx 2\pi (3/4) (k_S/k)^{4/3} \times (16/15),
$$

and

$$
J(k) \cong (8S_0/9\rho)^{4/3} 2\pi k^{-1/3} \times (8/5) (k_S/k)^{20/3}. \quad (51)
$$

So  $H(k_s/k)$  is  $8/5 (k_s/k)^{20/3}$ , compared with 1.2 for the case  $k \ll k_s$ . We see that the two expressions (47) and (51) for  $H(k_s/k)$  give almost identical values for  $k = k_s$ :

$$
H = 1.2
$$
 and 1.6, respectively.

We, therefore, conclude that the formulas hold up to (or down to) values rather close to that limit and that an interpolation should be easy.

Let us finally write down the cross-section formula corresponding to the two cases  $k \ll k_{s}$  and  $k \gg k_{s}$ .

According to (48) and (16) we have

$$
\sigma d\Omega = \frac{g^2}{8\lambda^4} (\rho_0/\gamma \rho_0)^2 (S_0/\rho)^{4/3} K^{-13/3} H(K/k_S) \sin^2 \chi d\Omega.
$$

Now  $(\rho_0/\gamma p_0) = (3/\gamma v_M^2)$ , where  $v_M^2$  is the mean square molecular velocity, and  $\gamma = 1.4$  for oxygen and nitrogen. Then finally, with  $K=2k \sin^1_2 \theta$ ,

$$
\sigma d\Omega \cong \frac{1}{2} g^2 (S_0^{4/3} v_M^{-4})
$$
  
 
$$
\times \lambda^{1/3} (2 \sin \frac{1}{2} \theta)^{-13/3} H\left(2 \frac{k}{k_S} \sin \frac{1}{2} \theta\right) \sin^2 \chi d\Omega.
$$

So, for the two cases, by using the extreme values (49) and (51) for H, and substituting  $S_0$  from (31), we finally get

$$
\sigma \cong 0.6g^2(v_0/v_M)^4 \lambda^{1/3} L_0^{-4/3} [2 \sin(\theta/2)]^{-13/3} \sin^2\chi,
$$
  
[2k sin(\theta/2)  $\ll k_S$ ]; (52)

$$
\sigma \cong 0.8g^2(v_0/v_M)^9 \lambda^7 L_0^{-3} L_M^{-5} [2 \sin(\theta/2)]^{-11} \sin^2\chi,
$$
  
2k sin(\theta/2)ggk\_S. (53)



FIG. 1. The integral  $J(k)$ .

## C. APPLICATION' TO SPECIAL CASES

## 1. Ionospheric Scattering

We apply our scattering formula (21) to the experi-We apply our scattering formula  $(21)$  to the experiment of Bailey *et al.*,<sup>12</sup> in which radiation of 50 Mc was scattered by the  $E$  layer. By using the same geometry scattered by the *E* layer. By using the same geometry<br>as Bailey *et al.*<sup>12</sup> we obtain for the ratio  $P_r/P_t$  of the power received to the power transmitted<br>  $\frac{P_r}{P} = 4\sigma$ .

$$
\frac{P_r}{P_t} = 4\sigma \frac{bA}{\sin(\theta/2)D^2}
$$

where  $b$  is the thickness of the scattering layer and  $A$ is the aperture of the receiving antenna. We use the following numbers which correspond to the setup in their experiments:

$$
\lambda = 10^2
$$
 cm,  $A = 3 \times 10^6$  cm<sup>2</sup>,  $D = 1.2 \times 10^8$  cm,  
 $b = 5 \times 10^5$  cm,  $\theta = 24^{\circ}$ .

For  $g^2$  we use (5) with  $\omega_N = 1.5$  Mc and the turbulence constants as given in (13a).

Substituting these values, we get<br> $P_r/P_t = 1.0 \times 10^{-18}$  [theoretic

$$
P_r/P_t = 1.0 \times 10^{-18} \text{ [theoretical from (21)],}
$$

whereas the experiments have given

 $P_r/P_i = 0.36 \times 10^{-18}$  (experimental).

This is a remarkably good agreement. We note that  $L = (2k \sin^3/2)$ <sup>-1</sup>=250 cm, which, according to (13b), is just equal to  $L_s$ . Hence, we actually should expect a somewhat smaller result than  $(21)$  and a steeper falling-off than  $(21)$  when going to larger angles or smaller wavelengths.

<sup>12</sup> D. K. Bailey et al., Phys. Rev. 86, 141 (1952).

## 2. Tropospheric Scattering

We have compared our results with the data of the experiments of (a) Round Hill, Massachusetts;<sup>13</sup> (b) Cheyenne Mountain, Colorado;<sup>14</sup> (c) the North Sea;<sup>15</sup> and (d) the Caribbean Sea.<sup>16</sup>

In experiment (a), we have (see Fig. 2):  $D=3\times10^5$ m;  $\lambda = 1.5$  cm;  $\alpha$ =antenna lobe width=0.6°;  $\theta$ =scattering angle=2.4°; and  $H=$ minimum height of intersection  $=1.8$  km. The useful scattering volume is  $V \cong (\frac{1}{2}D \sin \frac{1}{2}\alpha)^3/\sin \theta$ .

Throughout, in these four experiments, we shall compare the actual received (through scattering) power,  $P_{sc}$ , to the power received in free space over the same distance  $P_f$ . Thus the antenna characteristics are eliminated, and we get

$$
P_{sc}/P_f = 4(4\sigma V/D^2). \tag{54}
$$

Using (54) with (53), and with  $v_0$ ,  $L_0$  as given in (12), and  $g=0.45\rho$ , corresponding to the dry-air term of (6a), we get

$$
P_{sc}/P_f \cong 10^{-12}.
$$

The experimental data indicate a ratio of order  $10^{-8}$ , that is, a factor  $10<sup>4</sup>$  stronger.



FIG. 2. Geometry for tropospheric scattering (narrow beams).

<sup>13</sup> Private communication.

'4 Chambers, Herbstreit, and Norton, Nat. Bur. Standards, Report No. 1826, 1952 (unpublished). 's E. C. S. Megaw, Nature 166, 1100 (1950); Proc. Inst. Elec.

Engrs. III, 100, 1 (1953). "For these and similar experiments, see T. J. Carroll, Nat. Bur.

Standards, Report No. 1416, 1952 {unpublished).

In experiment (b) (and the following ones), we have a slightly diferent geometry: the beams are wide, and it is the form-factor of the scattering cross section rather than the lobe width which determines the useful scattering volume. We take, as a representative case in (b), a result obtained with  $D=3.6\times10^5$  m,  $\lambda=10$  cm,  $\theta$  = minimum scattering angle =  $3^{\circ}14'$ , and H = minimum height of intersection = 2.54 km, and get  $P_{se}/P_f$ = 10<sup>-11</sup>, compared to an observed  $10^{-7}$ . Again, we are a factor  $10^{-4}$  off.

That the Cheyenne mountain experiments are quite typical is seen from the data from (c) and (d). The very weak  $\lambda$  dependence of the cross section ( $\alpha \lambda^{1/3}$ ) make comparison easy, as far as orders of magnitudes go.

#### D. DISCUSSION

Further ionospheric experiments would be necessary to tell us whether the proposed picture of the dynamics of turbulence represents a reasonable approximation. Since in the troposphere the Reynolds numbers are generally much larger than in the ionosphere, we expect in general that our model is even better suited for tropospheric application. But there, in terms of dry-air fluctuations, an increase of the turbulent power output by a factor 10' would be necessary to account for the observed data, and this seems entirely inconceivable.

A natural explanation seems then to hold fluctuations in humidity responsible for the effect. If the assumptions made at the end of Sec. 2 are valid, we infer from (6a) and (14a) that, with say  $\rho' = 10^{-2} \rho$ ,  $T \approx 300^{\circ}$ , only a very small value for the large-scale humidity fluctuations  $(\Delta \rho')_0$  is necessary to explain the observed scattering:  $(\Delta \rho)_{0}^{2}/\rho'^{2} \cong 2 \times 10^{-4}$ . No measurements of  $(\Delta \rho')_0^2$  and the humidity fluctuation correlations are available yet. Measured large-scale humidity fluctuations in and outside cumulus clouds<sup>17</sup> give variations in  $\Delta \rho'/\rho'$  of the order of 10 percent over distances of 500 m. Fluctuations in a cloudless atmosphere are probably much smaller but might still be sufhcient (1 percent) to explain the effect.

A detailed investigation of the humidity efFect (and concurrently of the temperature effect) appears to be most important, and is under way.

We wish to thank Professor Henty Booker (Cornell) and Dr. J. de Bettencourt (M.I.T.) for many stimu lating discussions.

<sup>17</sup> J. S. Malkus, Sci. American 189, 31 (1953).